Changes in Stratigraphic Nomenclature by the U.S. Geological Survey, 1968

GEOLOGICAL SURVEY BULLETIN 1294-A







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By GEORGE V. COHEE, ROBERT G. BATES, and WILNA B. WRIGHT

CONTRIBUTIONS TO STRATIGRAPHY

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UNITED STATES DEPARTMENT OF THE INTERIOR

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CONTRIBUTIONS TO STRATIGRAPHY

CHANGES IN STRATIGRAPHIC NOMENCLATURE BY THE U.S. GEOLOGICAL SURVEY, 1968

By George V. Cohee, Robert G. Bates, and Wilna B. Wright

LISTING OF NOMENCLATURAL CHANGES

In the following table, stratigraphic names adopted, revised, reinstated, or abandoned are listed alphabetically. The age of the unit, the revision, and the area involved, along with the author's name and date of publication of the report, are given. The publications in which the changes in nomenclature were made are listed in the references at the end of this publication. The capitalization of age terms in the age column follows official usage.

One of the significant nomenclature changes in 1968 was the adoption by the U.S. Geological Survey of Holocene to replace Recent (Cohee, 1968). Holocene, meaning "wholly recent" and referring to the percentage of living organisms, originally was proposed as a "stage" following the Pleistocene "stage" by the Portuguese committee to the Third International Congress of 1885 (Morrison and others, 1957).

At the annual meeting of the American Commission on Stratigraphic Nomenclature, November 22, 1967, the Commission endorsed the use of Holocene instead of Recent and expressed the hope that the term "Holocene" would be adopted officially by various geological organizations. The formal term "Recent" was ambiguous in referring to sedimentary deposits, fossils, and present-day shells involving "Recent" or recent time. Holocene was given series rank equal to that of the Pleistocene because both vertebrate and invertebrate faunas reflect marked changes between Pleistocene and Holocene Epochs and the archaeological record provides means for subdividing Holocene deposits.

Changes in stratigraphic nomenclature

Revision and reference	Alibik Quartzite placed in the Menominee Group of the Anmikie Series. Are changed from Precambrian to	middle Precambrian. (Gair and Thaden, 1968.) - Alevea Linestone Member of Chugwater Formation raised in rank to Alevea Limestone of Chugwater Grandon.	(Pipiringos, 1998.) Altonian Substage (Frye and Willman (1960) adopted as basal substage of the Wisconsinan Stage. (Frye and others,	1968.) Wherever the Ames Limestone Member (of the Glenshaw Formation) can be recognized, the Conemangh is considered of group rank and consists of the Casselman and Glenshaw Formations and the Ames is considered that	top member of the Glenshaw Formation. (Reen and others, 1968.) Age changed from Pliocene to early Pliocene. (Kistler, 1968.)	- Age changed from Early and Middle Pennsylvanian to Late Mississippian and Early and Middle Pennsylvanian.	(Mallory, 1967.) Apache Canyon Formation of Tyrrell (1957) adopted. Overlies Willow Canyon Formation; underlies Shellenberger	Canyon Formation. (Finnell, this report, p. A32.) - Apsey Conglomerate Member adopted; overfies Hells Half Acre Tuff Member and underlies andesite of Tabbe	Mountain. (Krieger, 1968a.) Age changed from Eocene to Paleocene. (Hazel, 1968.)	44	Anocene. (Sato and Denson, 1907.) - Age changed from middle or late Tertiary to Oligocene.	(Steven and others, 1967.) Banner Limestone of Granger and others (1957) adopted as	Banner Formation. (Coats, 1989.) Bandstown Member adopted as middle member of Drakes Formation in north-central Kentucky. (Peterson, this	report, A36., Bathan Furnation adopted. (Stewart and McKee, 1968). Bathub Formation adopted (Drewes, 1968.) Bell Springs Member adopted as basal member of Nugget	Sandstone. (Pipiringos, 1968.) Age changed from Miocene and Pliocene(?) to late Miocene.	(Noble and others, 1968.). Bentley Formation extended into Texas. (Wilson, 1967).
Location	Michigan	Wyoming	Illinois and Wisconsin	Pennsylvania and Ohio	Nevada	Wyoming	Southeastern Arizona	Arizona	Maryland, Delaware, and	Arizona. Arizona. North Dakota, South Dakota,	Nebraska, and Wyoming. Colorado	Nevada	North-central Kentucky	Nevada. Southeastern Arizona. Wyoming.	Nevada	Texas.
Age	middle Precambrian	Triassic	Pleistocene	Late Pennsylvanian	early Pliocene	Late Mississippian and Early and Middle Penusylvanian.	Early Cretaceous	Miocene	Paleocene	Miocene early Miocene	Oligocene	Late Mississippian (Meramec). Nevada.	Late Ordovician (Cincin-natian).	Oligocene or Miocene Early Cretaceous. Triassic(?)	late Miocene	Pleistocene
Name	Ajibik Quartzite	Alcova Limestone Member (of Chugwater Formation).	Altonian Substage.	Ames Limestone Member (of Conemaugh For- Late Pennsylvanian mation).	Ammonia Tanks Member (of Timber Mountain Truft.)	Amsden Formation	Apache Canyon Formation (of Bisbee Group)	Apsey Conglomerate Member (of Galiuro Vol- Miocene canics).	Aquia Formation	Aravaipa Member (of Galiuro Volcanics)	Bachelor Mountain Rhyolite	Banner Formation	Bardstown Member (of Drakes Formation)	Battes Mountain Tuff	Belted Range Tuff	Bentley Formation

	CHANGES	IN	STRA	TIGR	APHIC 1	OME	NCLAT	URE		A3
In the Rio Grande embayment south of Frio River, the Bigford Member of the Mount Selman Formation is raised to formation rank, and the Mount Selman Formation is abandoned. (Earge, 1968.) Big Rock Conglomerate Member of Barker (1958) adopted. (Barker, this report, p. A21.) Birdo Creek Schist restricted. Some of the rocks formerly assigned to the Birch Creek Schist, restricted.	Keevy Feak Formation. (Wanringing, 1985.) Wherever the Ames Limestone Member (of the Glenshaw Formation) can be recognized, the Conemaugh is conconsidered of group rank and consists of the Caselman and Glenshaw Formations, and the Birmingham is considered the basal member of the Caselman Formation. (Roen and others, 1985.)	Group includes (secending order): Glance Conglomerate, Willow Canyon, Apache Canyon, Shellenberger Canyon, and Turney Ranch Formations (Finnell, this report, on A?8.)	Age changed from Eocene to Paleocene and Eocene. (Hazel, 1968.) Age changed from Oligocene(?) to early Eocene. (Steven	and others, 1907. Indian Mills Sandstone Member of Reger and Price (1928) adopted as member of Bluefield Formation. (Englund,	Bine Messa Tuff adopted. (Olson and others, 1968.) - Formation divided into (ascending order): Pride Shale, Glady Fork Sandstone, gray, red, Bramwell, and upper members. Age changed from Late Mississippian to Late Mississippian and Early Pennsylvanian. (Englund, Indeed)	Subsact, The Bossier is expanded to include all the dark-gray shale above the lower member of the Smackover Fornation and below the Schuler Formation. (Dickinson,	Bouse Formation adopted. (Metzger, 1968). Bramwell Member adopted. (Englund, 1968a.) Age changed from Silurian to Late Silurian. (Paylides and others, 1968.)	. Brightseat Formation made basal formation of Pamunkey Group. (Hazel, 1968.) - Age changed from Cretacous or early Tertiary to Late Cretacous (Harve and Dreaws 1968.)	. Wherever the Australianston Member (of the Glenshaw Pormation) can be recognized, the Commangh is considered of group rank and consists of the Casselman and Chenshaw Formations, and the Britsh Creek is considered a member of the Chenshaw Formation. (Roen	and others, 1968.) Brynt Draw adopted as basal member of Popo Agie Formation. (Pipringos, 1968.)
Southern Texas	. Pennsylvania and Ohio	DOUBLICAN AND THE COLOR	North Carolina and South Carolina. Southwestern Colorado	Virginia and West Virginia	Colorado Virginia and West Virginia	Texas, Louisiana, and Arkansas.	Southwestern Arizona Virginia and West Virginia New York	Maryland	- Pennsylvania and Ohio	. Wyoming
middle Eocene	Late Pennsylvanian	bally Olebaccous	Paleocene and Eoceneearly Eocene.	Late Mississippian	late Oligocene. Late Mississippian and Early Pennsylvanian.	Late Jurassic	Pliocene	PaleoceneLate Cretaceous	Late Pennsylvanian	. Late Triassic
Bigford Member (of Mount Selman Formation) - middle Eocene. Big Rock Conglomerate Member (of Kiawa Precambrian Mountain Formation).	Birmingham Shale Member (of Conemaugh Formation).	bisoee Group	Black Mingo FormationBlanco Basin Formation	Bluefield Formation	Blue Mess Tuff Bluestone Formation (of Pennington Group)	Bossier Formation (of Cotton Valley Group)	Bouse Formation Bramwell Member (of Bluestone Formation)—— Brayman Shale.	Brightseat FormationBronco Volcanics	Brush Creek Limestone Member (of Conemaugh Formation).	Brynt Draw Member (of Popo Agie Formation) Late Triassic

Name	Age	Location	Revision and reference
Buffalo Sandstone Member (of Conemaugh Formation).	Late Pennsylvanian	- Pennsylvania and Ohlo	Wherever the Ames Linestone Member (of the Glenshaw Formation) can be recognized, the Conemangh is considered of group rank and consists of the Casselman and Glenshaw Formations, and the Buffalo is considered a member of the Glenshaw Formation. (Roen and others)
Bull Ridge Member (of Madison Limestone) Bunker Andesite.	Early Mississippian (early Meramec). Oligocene.	West-central Wyoming	1968.) Bull Ridge adopted as upper member of the Madison Limestone. (Sando, 1968.) Name changed to Bunker Trachyandesite, age changed
Burnett Formation (of Puget Group)	Eocene	- Washington	from Eocene to Oligocene. (Steven and Epis, 1968.) Name abandoned. Lower part now assigned to Carbonado Formation, and upper part to Spiketon Formation.
Burns Formation (of Silverton Volcanic Group) Oligocene. Bursum Formation (of Magdalena Group) Early Per	Oligocene Early Permian.	. Colorado	(Gard, 1968.) Age changed from middle and late Tertiary to Oligoeene. (Luedke and Burbank, 1968.) Formation restricted geographically, to yicinity of Oscura
California Creek Member (of Totatlanika Schist)	Mississippian(?)	. Alaska	Mountains, central New Mexico. (Bachman, 1968.) California Creek Member adopted. (Wahrhaftig, 1968.)
Camarones Sandstone. Cambridge Limestone Member (of Conemaugh Formation).	Late Cretaceous. Late Pennsylvanian	. Puerto Rico	Camarones Sandstone adopted. (Pease, 1968a.) Wherever the Ames Linestone Member (of the Glenshaw Formation) can be recognized, the Comemangh is considered of group rank and consists of the Casselman and Glenshaw Formations, and the Cambridge is considered a member of the Glenshaw Formation. (Roen and others,
Campbell Mountain Member (of Bachelor	Early Devonian	. Kentucky, Tennessee, and Alabama. Colorado.	Age changed from Early and Middle Devonian to Early Devonian. (Botrot and Johnson, 1968.) Age changed from middle or late Tertiary to Oligocene.
Cancel Mayoure). Cancel Breeda. Cancel Hills Volcanics. Cantwell Formation.	Early to Late Cretaceous Late Triassic and Early Jurassic. Paleocene.	- Puerto Rico	(Steven and others, 1976, Cancel Precia and others, 1976, Cancel Precia adopted, (Pease, 1968a.) Age changed from Triassle and Jurassie to Late Triassic and Barly Jurassie, (Hayes and Drewes, 1988.) Age changed from Early Cretaceous to Paleocene. (Wolfe
Carbonado Formation (of Puget Group)	middle(?) Eocene	- Washington	and Wahrhaftig, this report, p. A41.) Carbonado Formation revised to include the Wilkeson Formation (abandoned) and the lower part of the Burnett Formatics (abandoned) and the lower part of the Burnett
Cardiff Conglomerate. Carpenter Ridge Tuff	Ordovician(?)Oligocenc	Maryland and Pennsylvania Colorado Puerto Rico	rounaton abandones, age changed non bocene to middle (7) Eocene (Gard, 1988.) Cardiff Conglomerate changed to Cardiff Metaconglomerate. (Southwise and Fisher, 1987.) Carpenter Ridge Tuff adopted. (Olson and others, 1968.) Carraizo Breccia adopted. (Pease, 1988.)
con close dan talto	Early of Middle Ordovician.	- Taano	Cash Creck Quartzite adopted. (Hobbs and others, 1968.)

Casper Formation	- Early, Middle, and Late Pennsylvanian and Early Permian.	Southern Wyoming	Age changed from Pennsylvanian and Permian to Early Pennsylvanian through Early Permian. (Mallory, 1967.)
Casselman Formation (of Conemangh Group)	Ä	- Pennsylvania and Ohio	Casselman Formation of Flint (1965) adopted wherever the Ames Linestone Member of the dribansw Formation) can be recognized. In those areas, the Conemaugh is considered a group and includes the Glenshaw and Casselman Formations. Members of the Casselman (in ascending order) are: Birmingham Shale, Elk Lick Limestone, Morgantown Sandstone, Clarksburg Limestone, Connellsville Sandstone, Summerhill Sandstone, and Elmand Chem.
Cerro Gordo Lava	Early to Late Cretaceous	Puerto Rico	ed as Cerro ortheastern
Chinquapin Metabasalt Member (of South Fork Mountain Schist). Chocolay Group (of Animikle Series)	Late(?) Cretaceous	Northwestern California C	
Chugwater Formation	. Triassic	. Wyoming.	(Gair and Thaden, 1968.) Chugwater Formation raised to group rank; includes (in ascending order): Red Peak Formation, Alcova Limessione, Crow Mountain Sandstone (locally Jelm Forma-
Chute Creek Member (of Totatlanika Schist) Cibao Formation	Mississippian(?). late Oligocene and early Miocene. middle Eocene.	Alaska	tion), and Popo Agie Formation. (Pipiringos, 1968.) Chute Creek Member adopted. (Wahrhaftig, 1968.) Age changed from Oligocene and Micoene to late Oligocene and early Micoene. (Eriggs, 1968.) Cook Mountain Formation in Texas restricted to eastern and central parts, replaced by Laredo Formation in Rio
Clarksburg Limestone Member (of Conemangh Late Pennsylvanian, Formation).	Miocene(?) and Pliocene Late Pennsylvanian	South-central Alaska	Grande emosyment, south of 1710 kiver. (Eargie, 1908.) (Clamgulchian Stage (Floral) adopted. (Wolfe and others, 1966.) Wherever the Ames Limestone Member (of the Glenshaw Formation) can be recognized, the Conemaugh is considered of group rank and consists of the Casselman and Glenshaw Formations, and the Clarksburg is considered consentant formations. And the Clarksburg is considered consentant formation.
Clayton Mine Quartzite. Clear Creek Formation. Clyde Formation.	Middle Ordovician or older Early Devontan Early Permian (Leonard)	Idaho	1968.) Clayfon Mine Quartzite adopted. (Hobbs and others, 1968.) Age changed from Early and Middle Devonian to Early Devonian. (Brouce and Johnson, 1968.) Age changed from Early Permian (Leonard): to Early Fermian (Leonard). (Myers, 1968.) In the southwesten part of the North Park area, the Middle Park Formation is reduced in rank and made the
Cockeysville Marble (of Glenarm Series)	late Precambrian(?)	Maryland, Pennsylvania, and Delaware.	basal member of the Commont Formation. (Hail, 1968.) Age changed from early Paleozoic(?) to late Precambrian(?) (Southwick and Fisher, 1967.) Compean Creek Gneiss adopted. (Gair and Thaden, 1968.)

Name	Age	Location	Revision and reference
Conemaugh Formation	Oligocene or olderLate Pennsylvanian	ColoradoPennsylvania and Ohlo	Age changed from middle or late Tertiary to Oligocene or older. (Olson and others, 1968.) Wherever the Armes Limestone Member (of the Glenshaw Formation) can be recognized, the Conemangh is considered of group rank and includes two formations, the Glenshaw (Passal) and Casselman Formations. The Armes to the two member of the Glenshow (Passal)
Conestoga Limestone	Early Ordovician	PennsylvaniaPennsylvania and Ohio	Annes is the out memor of the Clessaam, (Arota and Others, 1983.) Age changed from Late Cambrian and Early Ordovician to Early Ordovician. (Meisler and Becher, 1998.) Wherever the Ames Limestone Member (of the Clenshaw Formation) can be recognized, the Commangh is considered of group rank and consists of the Casselman and Glasniaw Formations, and the Comellsvilleis considered amender of the Casselman and Cambridge (Reconsidered Casselman Formation). (Reconsidered Casselman Formation). (Reconsidered Casselman Formation).
Continental Granodiorite	Precambrian middle Eocene Late Cretaceous and (or) early Tertiary. Early and Middle Silurian	Southeastern ArizonaTexas	1968.) Continental Granodiorite adopted. (Drewes, 1968.) Continental Granodiorite adopted. (Drewes, 1968.) Gook Mountain Formation in Texas restricted to eastern and central parts; replaced by Laredo Formation in Rio Grande embayment south of Frio River. (Eargle, 1968.) Age changed from Late(?) Cretaceous or Tertiary to Late Cretaceous and (or) early Tertiary. (Krieger, 1968b.) Lulbegrud Chay, Waco Lhrestone, and Estill Clay, all of Foerste (1908), adonted. Formation includes (in ascend-
Creede Formation Crow Mountain Sandstone Member (of Chugwater Formation). Darwin Sandstone Member (of Amsden Formation).	Oligocene	Colorado	ing order); Plum Creek, Oldham, Lulbegrud Shale, Waco, and Estill Shale Members. (Simmons, 1967.) (Steven and others, 1967.) Grow Mountain Sandstone Member raised to formation rank. Included in Chingwater Group. (Piphinges, 1968.) Age changed from Early Pennsylvanian to Late Mississip-
Devils Hole Formation. Dewitt Formation. Diamond Peak Formation.	late Bocene	South-central Colorado Eastern Texas	Parat. (Nature) 1997. Parat. (Nature) 1997. Epis, 1968. Point, 1968. Point, 1968. Pormation was named by Deussen (1914). In subsequent work by Deussen the name was not used nor was it ever used by any other author. Abandoned for nonuse. Age changed from Mississippian and Age changed from Mississippian and
Dillon Mesa Tuff. Dixville Formation. Drakes Formation.	Pennsylvanian. late Oligocene Middle Ordovician Late Ordovician (Cincinnatian).	Colorado	Early Pennsylvanian in the Carlin-Pinon Kange area. (Smith and Ketner, 1968.) Dilon Mesa Tuff adopted. (Olson and others, 1968.) Divylle Formation of Green (1964) adopted. (Harwood and Berry, 1967.) Bardstown Member adopted as middle member of Drakes Formation in north-central Kentucky. Overlies Rowland

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Member and underlies Saluda Dolomito Member. (Peterson, this report, p. A36.) Name changed from Eagle Valley Evaporite to Eagle Valley Formation. (Bartleson and others, 1983.) Elophant Head Quartz Monzonite adopted. (Drewes, 1968.) Wherever the Ames Limestone Moniber of the Glenshaw Formation) can be recognized, the Conemangh is considered of group rank and consists of the Casselman and Glenshaw Formations, and the Elk Lick is considered a manuper of the Casselman Pormation, and the Elk Lick is considered a	1998.) 1984. Ella Dounte adopted. (Hobbs and others, 1968.) Ellarsburg Formation extended from Washington into morth-central Oregon. (Waters, 1968.) El Octo Formation adopted. (Peass, 1968.) El Pico Clay adopted for the former upper unnamed member of the Mount Selman Formation (sbandoned) in the River (Brade embayment south of Frio River. (Bargie,	1998.) Buchantment Lake Formation adopted. (Gair and Thaden, 1908.) Bath Creek Glaciation adopted. (Sainsbury, 1967.) Bath Cry of Foerste (1908) adopted as Estill Shale Member. (Simmons, 1967.) Age changed from middle and late Tertiary to Oligocene. (Lucake and Burbant, 1908.) Wherever the Ames Limestone Member (of the Glenshaw Formation) can be recognized, the Conemangh is considered of group rank and consists of the Casselman and Glenshaw Formations, and the Ewing is considered a member of the Glenshaw Formations, and the Ewing is considered a member of the Glenshaw Formation. (Roen and others, 1988.)	. Fajardo Formation replaced by Rio Piedras Siltstone in Naranjito quadrangie, (Pease, 1965b.) - Falls Mills Sandstone Member of Reger and Price (1926) adopted. (Englund, 1963a.) - Age changed from Oligocene(?) to late Eocene. (Steven and Epis. 1983.) - Farmdale Silt of Frye and Willman (1960) adopted. (Frye and others, 1963.) - Farmdalina Substage of Frye and William (1960) adopted. Overlies Altonian Substage of Mrye and underlies Woodfordian Substage. (Frye and others, 1963.)	(Steven and others, 1937). Figurea Volcanies abandoned; replaced by Guaracanal Andestie, (Pease, 1968a.) Fish Caryon Tuff adopted. (Olson and others, 1968.) Age changed from middle or late Tertiary to Oligocene. (Steven and others, 1967.) Fleming Formation, previously abandoned, reinstated. (Anders and others, 1968.)
. Colorado	Idaho. Oregon. Puerto Rico. Southern Texas.	Michigan Alaska Kantucky. Colorado. Pennsylvania and Ohio.	Puerto Rico	
Pennsylvanian and Permian Late Cretaceous Late Pennsylvanian	Middle Ordovicianlate Micene and early Pilocene and early Pilocene Early to Late Cretaceousmiddle Eocene	middle Precambrian Holocene Middle Silurian Oligocene Late Pennsylvanian	Late Cretaceous Late Mississippian late Eocene late Pleistocene do Oligocene.	late Paleocene or early Eocene. late Oligocene
Eagle Valley Evaporito	Ella Dolomite Ellensburg Formation. El Ocho Formation El Pico Clay (of Claiborne Group)	Enchantment Lake Formation (of Chocolay middle Precambrian Group, Animike Series). Bisch Cleak Gladation. Bisch Cleak Gladat	Fajardo Formation Falls Mills Sandstone Member (of Hinton Formation), Farisita Conglomerate. Farmdale Silt. Farmdalian Substage (of Wisconsinan Stage) Farmers Creek Rhyolite.	Figuera Volcanies Fish Canyon Tuff Fisher Quartz Latite Fleming Formation

Neme	Аве	Location	Revision and reference
Ivanie	7450		
Fort Crittenden Formation	Late Cretaceous	Southeastern Arizona	Fort Crittenden Formation of Stoyanow (1949) adopted.
Fraction Tuff	middle Miocene	Nevada	(Drewes, 1968.) Age changed from late Miocene to middle Miocene. (Ander-
Frailes Formation	Late Cretaceous	Puerto Rico	son and Ekren, 1968.) The Frailes Formation, as used by Kaye (1959), is not a mappable unit in the report area and is included in the
			revised Guaynabo Formation, except for La Muda Lime- stone Member which is raised to formation rank. (Pease,
Franklinian Stage (Floral) Fultonian Stage (Floral) Galiuro Volcanics	early Eocene middle Eocene Miocene	West-central Washingtondo	Jaoos.) Frankinian Stage (Floral) adopted. (Wolfe, 1968.) Fultonian Stage (Floral) adopted. (Wolfe, 1968.) Includes (in ascending order): Holy Joe, Aravaipa, Hells
			Half Acre Tuff, and Apsey Conglomerate Members. Age changed from middle Tertiary or younger to Miocene. (Krieer, 1968b.)
Gallatin Group.	Late Cambrian	. Wyoming and Montana	Grove Creek Formation reduced to member rank and made upper member of Snowy Range Formation. Group in-
			cludes (in ascending order): Pilgrim Limestone and Snowy Range Formation. (Pierce and Nelson, 1988.)
Gardner Canyon Formation Ghost Rocks Formation Glady Fork Sandstone Member (of Bluestone	Triassic Paleocene and Eocene Late Mississippian	Southeastern Arizona Southern Alaska	Gardner Canyon Formation adopted. (Urewes, 1998.) Ghost Rocks Formation adopted. (Moore, 1969.) Glady Fork Sandstone Member of Reger and Price (1926)
Formation). Glenarm Series	late Precambrian(?)	_	adopted, (Englund, 1968a.) Age changed from early Paleozoic(?) to late Precambrian (?). In Maryland the following units are excluded from the
		v irginia.	Grenarm Series: Jamsvine Luymes, Mandue Sonne, Silver Run Linnestone, Urbana Phyllite, and Wakefield Morbler I. Morblerd the following units are included in
Glenshaw Formation (of Conemaugh Group) Late Pennsylvanian.	Late Pennsylvanian	Pennsylvania and Ohio	An Character Setters Formation, Cockeysville Marble, and the Wissalickon Formation (revised). (Southwick and Fisher, 1967.) Glonshaw Formation of Filint (1968) adopted wherever the Armes Limestone Member (of the Glenshaw Formation).
			can be recognized. In those areas, the Conemangh is considered a group and includes the Glenshaw and Casselman Formations. Members of the Glenshaw (in ascending order) are: Uffington Shale, Mahoning Saud-
			stone, Brush Creek Limestone, Buffalo Sandstone, Cambridge Limestone, Ewing Limestone, Saltsburg Sandstone, and Arnes Limestone. (Roen and others, 1998.)
Glory Hole Volcanics	Late Cretaceous and (or)	Arizona	Age changed from Late(?) Cretaceous to early Tertiary to Late Cretaceous and (or) early Tertiary. (Krieger, 1968b.)
Grape Creek Limestone Member (of Clyde Formation).	国	. Texas	Age changed from Early Permian (Leonard?) to Early Permian (Leonard). (Myers, 1968.)

•	CHANGES	IN STRATIG	RAPHIC NOM	ENCLATURE	A9
Gringo Gulel Volcanics adopted. (Drewes, 1968.) Grossman Romation adopted. (Cors. 1966.) Grosvenor Hills Volcanics adopted. (Cors. 1966.) Age changed from Miocene and Pliocene(?) to late Miocene (Noble and others, 1968.) Age changed from Miocene and Pliocene(?) to Miocene. (Kistler, 1968.) Grove Creek Formation reduced to member rank and miade upper, member. of Snowy Kange Formation.	(Priece and Nelson, 1997). (Unarcanal Andesite adoptical; replaces Figuera Volcanies which was abandoned, (Pease, 1968a.) Redefined to include parts of the Frailes Formation as used by Kaye (1959) and to exclude the younger La Muda Member. Includes Martin González Lava and Leprocomio Silstone Members. Age changed from Late Cretaconis(?) to Late Cretaconis. (Pease, 1968a.)	Hackett Sandstone Member of Reger and Price (1920) adopted. Englund, 1963a.) I vydell Sandstone Member adopted as member of Hance Hormation. (Englund, 1963b.) Hauser Lake Gneiss adopted, (Weis, 1963). Hells Half Acre Tuff Member adopted, underlies Apsey Conglomerate Member. (Krieger, 1963c). Age changed from middle and late Tertiary to Oligocene.	Under sur otherwish, 1908.) - Age charged from Pilocene(?) to late Tertiary. (Steven and others, 1967.) - Age charged from Pilocene(?) to late Tertiary. (Steven and others, 1967.) - Hinton Formation in Virginia and West Virginia divided into (in ascending order); stony dap Sandstone Member, ber, Hackett Sandstone Member, intide Stone Gap Member, Neal Sandstone Member, indide shale member, Tallery Sandstone Member, Pratter Shale Member, Falls, Mills, Sandstone Member, Pratter Shale Member, Falls, Mills, Sandstone Member, and upper shale mem-	Def. Chightin, 1962a. Def. Chightin, 1962b. Second and younger cpoch in the Quaternary Period. (Cothee, 1968). Holy Joe Mannber adopted. (Krieger, 1968b.). Holy Joe Mannber adopted. (Krieger, 1968b.). Horseshoe Shale Member adopted. (Mallory, 1967.). Age changed from middle or late Tertiary to Oligocene. (Sevon and others, 1967.) (Southwick and Fisher, 1967.) (Southwick and Fisher, 1967.) (Southwick and Fisher, 1967.) (Southwick and Fisher, 1967.) assigned to the Indian Hills Volcanics are now assigned assigned to the Indian Hills Volcanics are now assigned.	to the Orecu (Filter) Voicintes, (Anatrixon and Crease), 1967.) Indian Hills Sandstone Member of Reger and Price (1926) adopted. (England, 1968a).
Southeastern Arizona Novada Southeastern Arizona Novada Ovada Ovada Ovada Ovada	Puerto Ricodo	Virginia and West Virginia Tennossee Washington and Idaho Colurado	Maine Colorado	United States. Arizona. South-central Alaska. Wyouning. Colorado. Maryland.	Virginia_and West Virginia
Paleocene(?). Devonian or Mississippian late(?) Oligocene late Miocene do Late Cambrian	Paleocene	Lato Mississippian Middle Pennsylvanian Precambrian Miocene	Devonian (?) late Tertiary. Late Mississippian	Quaternary. Miocene Miocene and Pilocene(!). Late Mississippian and Early Pennsylvanian. Oligocene early Paleozoic(?) Precambrian.	Late Mississippian
Gringo Gulch Volcanies Grossman Formation. Grossman Formation. Grosseor Hils Volcanies Grouse Canyon Member (of Belted Range Tuff). late Miocene Grouse Canyon Member (of Indian Trail Formation). Grouse Canyon Member (of Indian Trail Formation). Grove Creek Formation.	Guaracanal Andesito	Hackett Sandstone Member (of Hinton Formation). Hance Formation (of Breathitt Group) Hauser Lake Gneiss Hells Half Acre Tuff Member (of Galiuro Volcanies). Henson Formation	Hildreths Formation Hinsdale Formation Hinton Formation	Holocene Epoch. Holy Joe Member (of Galiuro Volcanics) Honorain Siage (Floral). Horseshoe Shale Member (of Amsden Formation). Huerto Formation Ijamsville Phyllite Indian Hills Volcanics (of Alder Group).	Indian Mills Sandstone Member (of Fluefield Formation).

Name	Age	Location	Revision and reference
Indian Trail Formation	Miocene-	. Nevada	Age changed from Miocene and Pliocene(?) to Miocene.
Iron River Iron-Formation Member (of Michi-	Precambrian	. Michigan and Wisconsin	(Kistler, 1968.) Iron River Iron-Formation Member abandoned. (James
Ivydell Sandstone Member (of Hance Formation)	Middle Pennsylvanian	Northeastern Tennessee	and others, 1968.) - Ivydell Sandstone Member adopted. (Englund, 1968b.)
James Run Gneiss. James Run Gneiss. Jawone Conglomerate Member (of Kiawa	late Precambrian(?)	Maryland New Mexico	James Run Gneiss adopted. (Southwick and Fisher, 1967.) Jawbone Conglomerate Member of Barker (1958) adopted.
Mountain Formation). Jeffersonville Limestone	Early and Middle Devonian	Indiana and Kentucky	(Barker, this report, p. A21.) Age changed from Middle Devonian to Early and Middle
Jelm Formation	Late Triassic	. Wyoming	Devonian. (Boucot and Johnson, 1968.) John Formation assigned to the Chugwater Group; limited geographically to areas south of the Wind River basin. Members are Red Draw and Sine Creek. (Pipringos,
Jim Mountain Member (of Wapiti Formation)	99	Northwestern Wyoming	1968.) Jim Mountain Member adopted. (Nelson and Pierce, 1968.)
Jobos Formation	Eocene.	Puerto Rico	Tobos Formation adonted (Nelson, 1967.)
Josephine Canyon Diorite	Late Cretaceous. Late Jurassic.	Southeastern Arizona	Josephine Canyon Divide adopted. (Drewes, 1968.) Innerion Creek Sandstone reduced to member rank as
			Junction Creek Member of the Wanakah Formation.
Kane Wash Tuff.	Miocene	Nevada.	(Hansen, 1908.) Kane Wash Formation of Cook (1965) adopted as Kane
Kanouse Sandstone	Early Devonian	New York and New Jersey	Wash Tuff. (Noble, 1968.) Age changed from Early and Middle Devonian to Early
Keevy Peak Formation.	Precambrian or Paleozoic	Alaska	Devonian. (Boucot and Johnson, 1968.) Keevy Peak Formation adopted for rocks that were
Kiawa Mountain Formation	Precambrian	New Mexico	previously part of Birch Creek Schist. (Wahrhaftig, 1968.) Kiawa Mountain Formation of Barker (1958) adopted.
Kinnikinic Quartzite	Middle Ordovician	Idaho	(Barker, this report, p. A21.) Age changed from Late Ordovician to Middle Ordovician.
Kinnikinic Quartzite	-do	Central Idaho	
:			sorted quartzite of Middle Ordovician age in central Idaho. (Hobbs and others, 1968.)
Kodruda Ranch Complex Korna Dolomite	Cretaceous Tertiary or Cretaceous	Southern Alaska. Western Montana.	Kodiak Formation adopted. (Moore, 1969.) Kokoruda Ranch Complex adopted. (Sinedes, 1965.) Free Dalmite, in shood in the Cheoolay Cronn (Animitie
Kummerian Stage (Floral) La Garita Quartz Latite	early Oligocene	West-central Washington	Adua Joudhale and Planden, 1988. Kummerian Stage (Floral) adopted. (Wolfe, 1988.) Kan abancad from middle or late Pertiary to Olicoene.
Lake Fork Quartz Latite	Oligocene or older	Colorado	(Steven and others, 1967.) Name charged to Lake Fork Fornation; age changed from Micoene(?) to Oligoene or older. (Olson and others,
			1968.)

	CHAN	NGES I	N STRA	TIGRA	PHI	C NOM	ENCLATU	RE	AII
LaMuda Limestone Member raised to formation rank, La Muda Formation. (Peass, 1968a.) In the Rio Grande embayment south of the Frio River, the Laredo Formation replaces the Cook Mountain Formation which is restricted to eastern and central Texas north of the Frio River. (Eargle, 1968.) Leprocomio Limestone Member of Frailes Formation revised to Leprocomio Silistone Member of Guaynabo Formation. (Pease 1968.)	Age changed from Middle Ordovician to Middle and Late Ordovician. (Cressman, 1967).	Little Elk Granite of Taylor (1935) adopted. (Zartman and Stent, 1967.) Formation extended into northeastern Utah. (Sando, 1967.)	Table to the Intervol, in the 'type are only contained. Table Silurian(?) and Barly Devonian to Early Devonian. (Bondo and Johnson, 1988.) Loues Formation adopted. (Redden, 1968.) Loveland Loves changed to Loveland Silt in Illinois and Wisansein, (Free and others 1988.)	Lubbe Creek Formation adopted. (MacKevett, 1969.) Lubbegrud Shale Member of Foerste (1905) adopted. (Simmons, 1967.) Lyons Valley Member adopted. (Piptringos, 1968.)	McClure Mountain Complex adopted. (Shawe and Parker, 1967.)	Mackes d'Fanite adopted. (N'elson aur fross). Jobes.) Madere Ganyon Granodiorite adopted. (Drewes, 1988.) Bull Ridge adopted as upper member of the Madison Limestone. (Sando, 1988.) Madrid Formetion of Cariani (1989) adopted. (Osberg and	others, 1982. Wherever the Ames Limestone Member (of the Glenshaw Formation) can be recognized, the Concaught is considered of group rank and consists of the Gaselman and Glenshaw Formations, and the Mahoning is considered a member of the Glenshaw Formation. (Roen and others, 1968.)	Mainey Lava Member adopted. (Pease, 1968a.) Age changed from middle or late Tertiary to Oligocene. (Steven and others, 1967.) Semilla Sandstone Member adopted; overles Greenhorn Limestone Member and underles Juana Lopez Member	Z
Puerto Rico	Kentucky	South Dakota	New Hampshire South Dakota Wisconsin and Illinois	Alaska. Kentucky	. Colorado	Idaho Southestern Arizona West-central Wyoming Maine	. Pennsylvania and Ohio	. Puerto Rico	Maryland and Pennsylvania-
Late Cretaceousmiddle Eocene	Middle and Late Ordovician Kentucky early Precambrian Michigan.	Precambrian	Early Devonian Precambrian Pleistocene	Early Jurassic Middle Silurian Late Triassic	Precambrian or Cambrian	early Tertiary. Late Cretaceous. Early Mississippian (early Meramec). Silurian or Devontan.		Late CretaceousOligocene	early Paleozoic(?)
LaMuda Limestone Member (of Frailes Formation). Laredo Formation (of Claiborne Group) Leprocomio Limestone Member (of Frailes Formation).	Lexington Limestone	ange	Littletoù Fornation	Lubbe Creek Formation Lubegrud Shale Member (of Crab Orchard For- mation). Ivons Valley Member (of Popo Agle Forma-	tion). McClure Mountain Complex	Mackay Granite. Madera Canyon Granodiorite. Madison Limestone. Madrid Formation	Mahoning Sandstone Member (of Conemaugh Formation).	Mamey Lava Member (of Camarones Sandstone). Mammoth Mountain Rhyolite	Marburg Schist

Revision and reference	Martin González Lava Member adopted. (Pease, 1968a.) Mayflower Hill Formation adopted. (Osberg and others, 1968.) 1968.) Group in the Marquette area. (Galr and Thaden, 1968.) Meanard Quartzite placed in Cheoloya Group. Former age was Preeam Drian. Formation restricted to the relatively.	thick and massive vitreous quartitle. The lower part of the former Mesnard is placed in the Enchantment Lake Formation (new). (Gair and Thaden, 1968.) In the southwestern part of the North Park area, the Middle Park Formation is reduced in rank and made the basal member of the Coalmont Formation. (Hall, 1968.) Age changed from Devonian(?) and Early Mississippin to	Early Mississippian. (Sandberg and others, 1967.) Pioneer Sandstone of Glenn (1923) adopted as Pioneer Sandstone Member of Mingo Formation. (Englund, 1968b.) Age changed from Pennsylvanian and Permian to Early, Middle, and Late Pennsylvanian and Early Permian.		Mose Creek Member adopted. (Wahrhaftig, 1968.) Mose Creek Member adopted. (Wahrhaftig, 1968.) Middle Pennsylvanian. (Mallory, 1967.) Middle Pennsylvanian. (Mallory, 1967.) Wherever the Arnes Limestone Member of the Glenshaw Romation can be recognized, the Conemaugh is considered of group rank and consists of the Casselman and Glenshaw Formations, and the Morgantown is con-	sidered a member of the Casselman Formation. (Roen and others, 1988.) Morton Gnelss of Theil and Dutton (1935) adopted. (Stern and others, 1986.) Monthal (Effy Formation of Granger and others, 1988.) Monthain (Effy Formation of Granger and others (1957) adopted. (Coats, 1998.)
Location	Puerto Rico	. ColoradoIdaho.	. Tennessee	Michigan Puerto Rico Northeastern Ufah and southeastern Idaho. Texas and Louisiana.	Alaska. Wyoming. Pennsylvania and Ohio.	Minuesota
Age	Late Cretaceous	Palcocene	Middle Pennsylvanian Early, Middle, and Late Pennsylvanian and Early Permian	early Precambrian Late Cretaceous Mississippian Pleistocene	Mississippian (?) Barly and Middle Pennsylvanian. Late Pennsylvanian	Precambrian
Name	Markin González Lava Member (of Guaynabo Late Cretaceous. Formation). Mayflower Hill Formation. Menominee Group. Mesnard Quartzite (of Chocolay Group, Animdo	Middle Park Formation Milligen Formation	Mingo Formation (of Breathitt Group)	Mona Schist. Monacillo Formation. Monroe Canyon Limestone. Montgomery Formation.	Moose Creek Member (of Totatlanika Schist) Morgan Formation	Morton Gnelss. Morton Loess. Mountain City Formation.

	CH	ANGES	IN	ST	RAT	'IG	RA]	PHI	CI	IOMI	ENC	LA	TU	RE		A13
Mouti Selman Formation abandoned, and former members of formation raised to formation raised. In eastern and central Texas, they include (in ascending order): Reklaw Formation, Queen City Sand, and Weehes Formation. In the Rio Grande embayment south of the Frio River, they finclude (in ascending order): Biglord Formation and El Pico Clay (new). (Eargle, 1983.)	. Mount Wrightson Formation adopted. (Drewes, 1968.) . Mystic Creek Member adopted. (Wahrhaftig, 1968.) Age designated as early Paleozoic(?). (King and others,	ZZZ	pied by reason of use in Texas. Engund, 1903 Nelson Amphibolite of Granger and others (1957) adopted as Nelson Formation. (Coats, 1969.)	 Newman Lake Gneiss adopted. (Weis, 1968.) Pineville Sandstone Member of Hennen and Gawthrop (1915) adorted. (Fnefund. 1988a). 	. Nizha Moutfain Formation adopted. (MacKevett, 1969.) . Northcraft Formation placed in Puget Group. (Gard,	1903.) - Age changed from Cambrian (?) to Middle Ordovician (?).	(Pavlides and others, 190s.) Bell Springs Member of Triassic(?) age adopted as basal	member. Age demged from Early Jurassic to Triassic;) and Jurassic(?). (Pipiringos, 1963.) Age changed from Pleistocene(?) to Pleistocene. (Soister,	1967.) - Age changed from Eocene(?) to middle Eocene. (Yates	and Engels, 1983.) Age changed from late Eocene to Oligocene. (Wolfe, 1983.) Ortega Quartzite of Barker (1958) adopted. (Barker, this	report, p. A21.) Age changed from middle or late Tertiary to Oligocene. (Stevan and others 1937.)	- Age changed from Miocene(?) and Pliocene to late Miocene. (Kistler, 1968.)	- Age changed from Miocene (?) and Pliocene to late Miocene.	Pajaros Tuff abouted. (Pease, 1963a). Age changed from middle and late Miocene to Pliocene and	Prightseat Formation added as basal formation of group; age charged from Eocene to Paleocene and Eocene. The Add the Committee of the Eocene and Eocene.	Tantano Pornation of Tolman as defined by Brennan (1962) adopted. (Finnell, this report, p. A35.)
. Техаѕ.	Alaska. North Carolina, Tennessee,	Puerto Rico Southern Alaska Virginia and West Virginia	. Nevada	- Eastern Washington - Pennsylvania, Virginia, and West Virginia.	Alaska	Maine	Wyoming	Colorado	Northeastern Washington	Washington New Mexico	- Colorado	Nevada	do	Puerto Rico Southeastern California	Maryland, Delaware, and Virginia.	Southeastern Arizona Maryland and Virginia
middle Eocene	Triassio	Paleocene Miocene Late Mississippian	op-	Precambrian Early Pennsylvanian	Middle and Late Jurassic	Middle Ordovician(?)	Triassic(?) and Jurassic(?)	Pleistocene	middle Eocene	Oligocene Precambrian	Oligocene	late Miocene	-do	Early to Late Cretaceous	. Paleocene and Eocene	early Oligocene to early Miocene. Paleocene
Mount Selman Formation (of Claiborne Group). middle Eocene.	Mount Wrightson Formation Mystic CreekMember (of Totatlanika Schist) Nantahala Slate	Naranjito Fornation	Nelson Formation	Newman Lake Gneiss New River Formation	NizinaMountain Formation	North Haven Greenstone	Nugget Sandstone	Nussbaum Alluvium	O'Brien Creek Formation	Ohanapecosh Formation Ortega Quartzite	Outlet Tunnel Member (of La Garita Quartz	pah Canyon Member (of Paintbrush Tuff)	Paintbrush Tuff (of Piapi Canyon Group)	Pájaros Tuff. Palm Spring Formation.	Pamunkey Group.	Pantano Formation

Name Penters Chert	Age Early Devonian	Location	Revision and reference Age changed from Early or Middle Devonian to Early
Peoria Loess	Pleistocene	Illinois and Wisconsin	Devonian. (Boucot and Johnson, 1988.) Peoria Loess of Frye and Willman (1960) adopted. (Frye
Perry Mountain Formation	Silurian(?)	Maine	and others, (1968).) Perry Mountain Formation of Cariani (1959) adopted.
Peters Creek Quartzite	late Precambrian(?)	Virginia and Pennsylvania	(Osberg and others, 1968.) Peters Creek Quartzite abandoned in Maryland; remains in good usage in Pennsylvania and Virginia. (Southwick
Phoenix Park Member (of La Garita Quartz Latite).	Oligocene	. Colorado	and Fisher, 190.) Age changed from middle or late Tertiary to Oligocene.
Piapi Canyon Group	late Miocene and early Pliocene.	Nevada	Age changed from Miocene (?) and Pliocene to late Miocene and early Pliocene. (Kistler, 1968.)
Fray une Formation Piña Siltstone Member (of El Ocho Formation) Pine Butte Member (of Sundance Formation)	Ongocene Early to Late Cretaceous Late Jurassic	Colorado Puerto Rico Wyoming	Age changed from middle and late Tertiary to Oligocene. (Luedke and Burbank, 1988.) Piña Slitstone Member adopted. (Pease, 1968a). Town: sanderne, bade rannored from Badmeter Shalo
Pineville Sandstone Member (of New River	Early Pennsylvanian	Virginia and West Virginia	Member and named Pine Butte Member. (Pipringos, 1965.) Pineville Sandstone Member of Hennen and Gawthron
Formation). Finner Sandstone Member (of Mingo Forma-	Middle Pennsylvanian	Tennessee	(1915) adopted. (England, 1968a.) Pioneer Sandstone Member of Glenn (1925) adopted.
Juon). Piper Gulch Monzonite. Piscataway Indurated Marl Member (of Aquia Rormation)	Triassic. Paleocene	Southeastern Arizona	(Englund, 1968). Piper Gulch Monzonite adopted. (Drewes, 1968.) Age changed from early Eocene to Paleocene. (Hazel, 1968.)
Pitis Meadow Granodiorite	Precambrian Late Triassic.	Colorado	Pitts Meadow Granodiorite adopted. (Hansen, 1968.) Raised in rank to Popo Agle Formation of Chugwater Group. Formation redefined and restricted to analeme-
Poxono Island Formation	Late Silurian	stern Pennsylvania, western New Jersey, utheastern New	rich beds of Keller (1952), (Pipiringos, 1968.) Poxono Island Shale of White (1882) adopted as Poxono Island Formation. (Epstein and others, 1967.)
Pratter Shale Member (of Hinton Formation) Pride Shale Member (of Bluestone Formation)	Late Mississippiando	Virginia and West Virginia	Pratter Shale Member adopted. (Englund, 1968a.) Pride Shale Member of Reger and <i>Frice</i> (1926) adopted.
Pringle Andesite	Oligocene	Southern Colorado	(Englund, 1968a.) Name changed from Pringle Andesite to Pringle Latite; age clanged from Perifary to Oligocene. (Steven and
Puget Group	Eocene and Oligocene(?)	Washington	Epis, 1968.) Pierce County the revised Puget Group includes (in ascending order): the Carbonado, Northeraft, and Spike-
Do	early Eocene to early Oligo- cene.		ton tomations, the witteen and burnet, romations, formerly in Puget Group, are abandoned. (Gard, 1968.) In King County the age of the Puget Group is changed from Eocene and Oligocene(?) to (1) early Eocene to

CHANGES IN STRATIGRAPHIC NOMENCLATURE	A15
early Oligocene in the Green River area, and (2) middle Eccene to early Oligocene in the Tiger Mountain area. (Wolfe, 1963.) Age changed from Pennsylvanian to Middle Pennsylvanian (Des Moines). (Mallory, 1967.) Age changed from Endlory, 1967.) Age changed from middle and late(?) Eccene to middle Eccene. (Wolfe, 1968.) Age changed from Pilocene to early Pilocene. (Kistler, 1968.) Age changed from Filocene to early Pilocene. (Kistler, 1968.) Ranchesterj Linnestone Member adopted. (Mallory, 1967.) Ranchesterj Linnestone Member adopted. (Mallory, 1967.) Ranchesterj Linnestone Member adopted. (Wolfe, 1968.) Rayerian Stage (Floral) adopted. (Wolfe, 1968.) Recent Epcoh abandoned and replaced by Holocene Epoch as second and younger epoch of the Quaternary Period. (Cohee, 1968.) Rech Draw Member adopted. (Pipiringos, 1968.) Raised to formation rank, Red Peak Formation of Chug-water Group. (Pipiringos, 1968.) Raised to formation rank, Red Peak Formation of Chug-water Group. (Pipiringos, 1968.) Raised to formation rank, Red Peak Formation raised to formation rank. Red Peak Formation raised to formation rank. Relaw Formation, Mount Selman Formation abandoned. (Eargle, 1968.) Relation abandoned. (Eargle, 1968.) Relation abandoned. (Eargle, 1968.) Recent Mannber and maned Pine Butte Member. (Pipiringos, 1968.) Recentaged from Itale Coccene and Oligocene(?) to late Eocene and early Oligocene. (Wolfe, 1968.) Richland Loess of Frye and Wilman (1969) adopted. (Pease, 1968.) Recet wisconsinan: (Frye and others, 1969.) Richland Loess of Frye and Wilman (1969) adopted. (Pease, 1968.) Recet Wisconsinan: (Frye and others, 1968.) Recet wanged from Paleocene to Picistocene. (Steven and others, 1967.) Rocket Wash Member adopted. (Roska, 1968.) Rocket Wash Member adopted. (Roska Formations adopted. (Roska Formations adopted.) Rocket Wash Romber adopted. (Roska Formation adopted.) R	Roxana Silt of Frye and Willman (1960) adopted as basal unit of Wisconsinan Stage. Age is early Wisconsinan (Altonian). (Frye and others, 1968.)
Wyonning Texas. Washington Woyoning Wyonning Wyonning Wyonning Ado. do. Texas. Texas. Washington Colorado. Go. do. do. Abashington Washington Washington Washington Washington Washington Washington	. Illinois and Wisconsin
Middle Pennsylvanian (Des Moines). do do early Pliocene. Gravian. Silurian(?). late Eocene Quaternary. Late Triassic Early Triassic Barly Triassic Late Jurassic Iate Eocene and early Oligocene. Pleistocene. Late Cretaceous. Late Cretaceous. Late Jurassic late Jurassic middle Eocene and Eocene(?). Pleistocene. Late Gravianian(?). Late Jurassic Late Jurassic Late Cretaceous. Late Jurassic Late Jurassic Indidle Eocene middle Eocene Late Cretaceous. Late Jurassic Late Jurassic Late Jurassic Late Jurassic Late Jurassic Late Jurassic	late Pleistocene
Quadrant Quartzite	Roxana Silt.

Name	Age	Location	Revision and reference
Salero Formation Salisaw Formation Saltsburg Sandstone Member (of Conemaugh Formation).	Late Cretaceous. Barly Devonian. Late Pennsylvanian	Southeastern ArizonaOklahomaPennsylvania and Ohio	Salero Formation adopted. (Drewes, 1968.) Age changed from Early or Middle Devonian to Early Devonian. (Boucot and Johnson, 1968.) Wherever the Ames Limestone Member (of the Glenshaw Formation) can be recognized, the Conemaugh is considered of group rank and consists of the Casselman and
San Juan Formation	Oligocene or oldermiddle Bocene	ColoradoNortheastern Washington	Glenshaw Formations, and the Saltsburg is considered a member of the Glenshaw Formation. (Roen and others, 1968.) Age changed from middle or late Tertiary to Oligocene or older. (Olson and others, 1968.) Age changed from Eocene(?) to middle Eocene. (Yates and
San Sebastián Formation (of Río Guatemala Group). Santa Olaya Laya. Sapinero Mesa Tuff.	middle Oligocene	Puerto Rico	Engels, 1968.) Age changed from Oligocene to middle Oligocene. (Briggs, 1968.) Santa Olaya Laya of Lidiak (1965) adopted. (Pease, 1968.) Sapinero Mesa Tuff adopted. (Olson and others, 1968.)
Schoharie Grit	younger. Early Devonian		Age changed from Late Ordovician to Middle Ordovician Ordovician Ordovician (Ruppel, 1968.) Name changed to Schoharie Formation. (Boucot and Johnson, 1968.)
Sciute: Formation (of Cotton vaney Group) Late Justissic. Seldovian Stage (Floral)	Late Jurassic	Lexas, Louisiana, and Arkansas. South-central Alaska	Scilluer formation redemined to incute an rocks between the top of the Buckner Formation and the base of the Cretaceous rocks, except for the Q tongue of the Bossier Formation. (Dickinson, 1968.) Seldovian Stage (Floral) adopted. (Wolfe and others, 1966.) Semilia, Sandstone Member adopted. (Dane and others,
Setters Formation (of Glenarm Series)	late Precambrian(?)	Maryland, Pennsylvania, and Delaware.	1968.). Age changed from early Paleozoic(?) to late Precambrian(?). (Southwisk and Fisher, 1967.) Age changed from middle or late Tertiary to Oligocene.
Sheep Creek Member (of Totatlanika Schist) Shellenberger Canyon Formation (of Bisbee Group).	Mississippian(?)	Central Alaska	(Skeva and others, 1987). Sheep Creek Member adopted. (Wahrhaftig, 1968.) Shellenberger Canyon Formation of Tyrrell (1957) adopted. (Finnell, this report, p. A32.) Age changed from Tertlary to Rocene. (Yates and Engels,
Siamo Slate	middle Precambrian		1968.) Placed in Menominee Group of Animikie Series. Age changed from Precambrian to middle Precambrian. (Gair and Thaden, 1962).
Silver Run Limestone	early and initiate rocene early Paleozoic(?)	Western Oregon	Nathe Changer (run Sucker Arver vocanic series to Sucker River Volcanics. (Snayely and others, 1998.) Silver Run Limestone is excluded from Glenarm Series. (Southwick and Fisher, 1967.) Age changed from middle and late Terttary to Oligocene. (Luedke and Burbank, 1998.)

CHANGES IN SIGNIFICATION	10 1101121101111 01111
Sips Creek Member adopted, (Pipiringos, 1968.) Sikralidak Formation adopted, (Moore, 1969.) Sikralidak Formation adopted, (Moore, 1969.) Sinalis Falls Formation of Furiong (1960) adopted. (Usberg and others; 1963.) Age changed from middle or late Tertiary to Oligocene. (Skeven and others; 1967.) Grove Creek Formation reduced to member rank and made upper member of Snowy Range Formation. (Fierce and Nolson, 1968.) South Fork Mountain Schist adopted. (Blake and others, 1967.) Spiketon Formation adopted. (Gard, 1968.) Spiketon Formation adopted. (Gard, 1968.) Spiketon Formation adopted. (Drewes, 1968.) Age changed from Middle and Late Triassic to Early to Lafe Prissics. (Silberling, 1968.) Age changed from Middle and Late Triassic to Early to Lage changed from Middle and Swanson, 1968.) Age changed from Middle and Late Triassic to Late Originated from Middle Pilocene to middle(?) and late Member of Morene, Wisser, 1968.) Age changed from Cretaceous or Tertiary to Late Cretaceous. (Hayes and Drewes, 1968.) Age changed from Cretaceous or Tertiary to Late Cretaceous. (Hayes and Drewes, 1968.) Age changed from Cretaceous or Tertiary to Late Cretaceous. (Hayes and Drewes, 1968.) Wilther 1968.) Wherever the Ames Limestone Member (of the Glenshaw Formation) can be recognized, the Conemaugh is considered of group rank and consists of the Casselman and Glenshaw Formations, and the Summerhill is considered a member of the Casselman Formation. (Roen and	Outners, 1993. Outners, 1998. Sykesville Formation abandoned in Maryland. Rocks are now included in boulder guelss lithoideces of Wissahicken Formation. Sykesville Formation abandoned in Maryland. Rocks are now included in boulder guelss lithoideces of Wissahicken Formation. Sykesville Formation still in good usage in Pennsylvania. (Southwick and Fisher, 1967.) Age changed from Edily Pennian (Leonardy to Early Permian (Leonardy to Early Permian (Leonard). (Myers, 1968.) Age changed from Oilgocene(?) to early Eocene. (Steven and others, 1967.) Temporal Formation adopted. (Drewes, 1968.) Age changed from middle(?) and late Eocene to middle and late Eocene. (Wolfe, 1968.) Age changed from Midocene to early Pilocene. (Kistler, 1968.) Age changed from Miocene(?) and Pilocene to late Miocene. (Kistler, 1968.) Age changed from Miocene(?) and Pilocene to late Miocene. (Fistler, 1968.)
Wyoming South-central Alaska Johnande Colorado Wyoming and Montana Northwestern California Southeastern Arizona Nevada Oregon Nevada Southeastern Arizona Nevada Coregon Nevada Southeastern Arizona - Central Kentucky	Wyoming
Late Triassic Bocone Silurian(?). Oligocene Late Cambrian Late(?) Cretaceous. Inassic. Barly to Late Triassic. middle(?) and late Miocene late Miocene Late Cretaceous Late Cretaceous Late Pennsylvanian.	r (of Hinton Forma- Late Mississipplan. of Clyde Formation). Early Permian (Leonard) early Bocene (of Puget Group) Barly Cretaceous indidle and late Eocene (of Piapi Canyon early Pliocene
Sips Cireck Member (of Jelm Formation) Sikidalidak Formation Sindalidak Formation Sinalis Falls Formation Snowle Mountain Quartz Latite Snowy Range Formation (of Puget Group) Squaw Gulch Granite Star Peak Formation. Steens Basalt Stockade Wash Member (of Paintbrush Tuff) Sugarloaf Quartz Latite Sugarloaf Quartz Latite Sugarloaf Quartz Latite Summerhill Sandstone Member (of Conemaugh Formation).	Sykesville Formation

Revision and reference	Age changed from Miocene (?) and Pliocene to late Miocene. (Kistler, 1968.) Stratigraphically restricted to exclude rocks assigned to Reevy Peak Formation. Subdivided into five nearly manded members (in ascending order): Moose Creek, Chalfornia Creek, Chatte Creek, Mystic Creek, and Sheep Oreak. (Wahrhattig, 1968.) Age changed from middle or late 'Tertiary to Oligocene.' (Steven and others, 1967.) Extended to Idaho. (Sandberg and others, 1967.) Trout Peak Trachyandesite adopted. (Nelson and Pierce, 1968.) Truillo Alto Linnestone reduced in rank and made member of Monacillo Formation. (Pease, 1968.) Age changed from Miocene and Phocene(?) to late Miocene. (Kistler, 1968.) Tugidak Formation adopted. (Moore, 1969.) Tugidak Formation adopted. (Moore, 1969.) Age changed from Jae Ecoene to late Ecoene and early Oligocene. (Wolfe, 1968.)	Turney Ranch Formation of Tyrrell (1957) adopted. Top formation of Bisbee Group. (Finnell, this report, p. A33). Twocreekan Substage of Frye and Willman (1960) adopted as second youngest substage of Wisconsinan Stage. Overlies Woodfordian Substage and underlies Valderan Substage. (Frye and others, 1988.) Wherever the Ames Linestone Member (of the Glenshaw Formation) can be recognized, the Concamarth is considered a member of group rank and consists of the Casselman and Glenshaw Formations, and the Uffington is considered a member of the Glenshaw Formation. (Roen and others, 1968.) Age changed from Middle and Late Ordovician to Middle Ordovician. (Bachman, 1968.) Urbana Phyllite excluded from Glenarm Series. (Southwick and Fisher, 1967.) Urbana Phyllite excluded from Glenarm Series. (Southwick and Fisher, 1967.) Valderan Substage of Frye and Willman (1960) adopted as uppermost substage of Wisconsinan Stage. (Frye and others, 1968.)
Location	Nevada Central Alaska Colorado Montana, Wyoming, and Idaho. Wyoming Puerto Rico Nevada do South-central Alaska	South Dakota
Age	Mississippian (?) Oligocene Late Devonian late Eocene Late Cretaceous Late Miocene Miocene Pliocene Pliocene Pliocene Pliocene late Eocene and early Oligocene	Early Cretaceouslate Pleistocene Latc Pennsylvanian Middle Ordovician early Paleozoic(?) Triassic
Name	Topopah Spring Member (of Paintbrush Tuff) late Miocene Totatlanika Schist	Turney Ranch Formation (of Bisbee Group) Barly Cretaceous Twocreekan Substage

iocene to early	(1905) adopted. (Simmons	o late Miocene.	Series. (South-	Late Triassic.	nk to Junction	(Hansen, 1968.) erce, 1968.)	y to Oligocene.	Smith (1925), arg and others,	ner, 1968.) mian to Middle (Mallory, 1967.) nan Formation) Mount Selman	(1960) adopted.	6.)	or older. (Olson	ounger to Late 168.) Animikie Series. e Precambrian.	assigned to the	Fertiary to Late ger, 1968b.) (1957) adopted.	y to Oligocene. If the Glenshaw naugh is consid- Casselman and e is considered n. (Roen and
Pliocene to late M	rste (1905) adopt	ne and Pliocene t	ed from Glenarm	ssic or Jurassic to	ne reduced in ra	akan Formation. ed. (Nelson and Pi	le or late Tertiar	d by Perkins and Formation. (Osb	1. (Smith and Ket sylvanian and Per an and Permian. oer (of Mount Selrons, Formation;	rye and Willman (Woodfordian).	ptcd. (Hernon, 196	ne(?) to Oligocene	nsylvanian or yough. (Feininger, 1) hocolay Group, Asambrian to midd	8.) andoned; strata (revised) (Gard	cretaceous or riy Tertiary. (Kright)	1.4.21.7 11e or late Tertian 67.7 1.6.2 for the Concr 1.6.2 for the Concr 1.6.2 for the Concr 1.6.2 for the Concr 2.6.3 for the Wilmon Sschman Formatic
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late Miocene to early Pliocene.	Middle Silurian	late Miocene	early Paleozoic(?)	Late Triassic	Late Jurassic	early(?) and middle(?)	Eocene. Oligocene	Silurian(?)	Early Mississippian	late Pleistocene	Precambrian	Oligocene or older	Late Pennsylvanian or younger. middle Precambrian	Eocene	Late Cretaccous and (or) early Tertiary. Early Cretaceous	Oligocene
Waccamaw Formation	Waco Member (of Crab Orchard Formation) 1	Wahmonie Formation1	Wakefield Marblee	Walnut Gap Volcanics	Wanakah Formation	Wapiti Formation 6	Wason Park Rhyolite	Waterville Formation	Webb Formation. Weber Sandstone. Weches Greensand Member (of Mount Selman Pormation).	Wedron Formation	Wehutty Formation	West Elk Breccia	Westerly Granite	Wilkeson Formation (of Puget Group)	Williamson Canyon VolcanicsWillow Canyon Formation (of Bisbee Group)	Willow Creek Member (of Bachelor Mountain Rhyolite). Wilmore Sandstone Member (of Conemaugh Formation).

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Name	Windy Gulch Member (of Bachelor Mountain Oligocene		Wisconsinan Stage	Wissahickon Formation (of Glenarm Series) ls	Woodfordian Substagela	Woodruff Formation	Yucca Mountain Member (of Paintbrush Tuff) late Miocene.	Zooks Corner Formation C

ORTEGA QUARTZITE AND THE BIG ROCK AND JAWBONE CONGLOMERATE MEMBERS OF THE KIAWA MOUNTAIN FORMATION, TUSAS MOUNTAINS, NEW MEXICO

By FRED BARKER

The Tusas Mountains of northern New Mexico contain extensive exposures of Precambrian quartzite and conglomerate. In a reconnaissance investigation Just (1937, p. 43, and pl. III) applied the name Ortega Quartzite to these rocks for their exposures from Ojo Caliente to Jawbone Mountain (T. 29 N., R. 6 E.) with the type locality in the Ortega Mountains. In a more recent study in the Las Tablas, Canon Plaza, Burned Mountain, and Mule Canvon 71/2-minute quadrangles, Barker (1954, 1958) subdivided Just's Ortega Quartzite. In conformance with this work, the name Ortega Quartzite is applied only to the 14,000- to 20,000-foot-thick, lower part of this very thick section of quartzite and conglomerate as exposed in the center of section 25, T. 27 N., R. 8 E., and to the west and southwest. The base of the Ortega Quartzite is not exposed in the Tusas Mountains. The strata overlying the redefined Ortega Quartzite are named the Kiawa Mountain Formation, after the exposures at Kiawa Mountain, the type locality (secs. 3-5, 8-10, T. 27 N., R. 6 E.; Barker, 1958). The uppermost part of the Kiawa Mountain Formation is not preserved in the Tusas Mountains.

The lowest unit of the Kiawa Mountain Formation in the Las Tablas quadrangle is a pebble conglomerate and is named the Big Rock Conglomerate Member (Barker, 1958) for exposures about 0.2 to 1 mile north and northwest of Big Rock. The type locality is in the SE¼ sec. 27, T. 27 N., R. 8 E. About 15 miles to the northwest, in the Burned Mountain quadrangle, the basal part of the Kiawa Mountain Formation is pebble conglomerate, fine-grained conglomerate, and quartzite and is named the Jawbone Conglomerate Member (Barker, 1958). The type locality is in the N½ sec. 19, T. 29 N., R. 7 E. Excellent exposures of this member are found at Jawbone Mountain which lies in the northwest corner of the Burned Mountain quadrangle and in the northeast corner of the Cebolla quadrangle. The stratigraphic relations of these two conglomerates are not known; however, there is a possibility that the Jawbone is younger than the Big Rock (Barker, 1958, pl. 1).

The Ortega Quartzite, restricted, and the Kiawa Mountain Formation and its two members, the Big Rock and Jawbone Conglomerate

Members, are herein adopted.

The Kiawa Mountain Formation consists principally of quartzite and contains some units of conglomerate and amphibolite. The quartzite layers are light blue to gray, dense, vitreous, and typically very fine grained to medium grained. Layers containing granules and

pebbles of quartz and, less commonly, of ferruginous chert or ironformation occur throughout. At Kiawa Mountain and to the northwest, the quartzite consists of 90 to 98 percent quartz with minor amounts of hematite, kyanite, manganian and alusite, rutile, and other minerals. Northeast of Kiawa Mountain, in the canyon of Spring Creek, the quartzite is feldspathic and muscovitic. South of Kiawa Mountain, on La Jarita Mesa, the quartzite of this formation grades into the Petaca Schist of Just (1937, p. 43), which is a muscovitic quartzose schist. The Petaca Schist is only found adjacent to the pegmatite bodies of the Petaca District, and Just ascribes the muscovite in the schist to emanations from the pegmatites.

The quartzite layers of the Kiawa Mountain Formation are intensively folded, and so their thickness cannot be measured accurately. However, they are estimated to be 5,000 to 10,000 feet thick (Barker,

1958, p. 10).

The Big Rock Conglomerate Member consists of interlayered gray quartzose conglomerate and pebbly quartzite. The clasts are of quartz and jasper, and they generally range from ½ to 5 inches in maximum dimension. The matrix of both rock types typically is 1/16 to 4 mm (millimeters) in grain size and shows a mosaic fabric. Quartz forms 60 to 80 percent; microcline and muscovite are next in abundance; and biotite, garnet, and hematite are accessory in the matrix. Both current bedding and cross-stratification are common in the Big Rock Conglomerate Member. This member is intricately folded. and is estimated to be 100 to 200 feet thick near Big Rock. It pinches out to the southeast. To the northwest it also apparently pinches out: here, however, this conglomerate was intruded by Precambrian rhvolite and later was partly covered by Tertiary conglomerate (Barker, 1958, pl. 1), so its stratigraphic relations are obscured.

The Jawbone Conglomerate Member is comprised of interlayered, gray to light gray, vitreous, massive, quartzose conglomerate. quartzite, and pebbly conglomerate. Granules and well-rounded pebbles of quartz and jasper of 1/4 to 1 inch in grain size form the coarsest part of the conglomerates. The matrix of the conglomerate and the quartzite consists of recrystallized quartz, mostly 1/8 to 1 mm in grain size, and accessory kyanite, hematite, rutile, and muscovite. Bedding and cross-stratification may be seen in many outcrops of the Jawbone. Because this member is folded, its thickness may not be directly measured; however, it probably is at least 500 feet thick and may be more than 2,000 feet thick.

REASONS FOR ABANDONMENT OF THE PORTAGE GROUP By Wallace de Witt, Jr.

In our study of the stratigraphy of the Upper Devonian rocks in western and west-central New York (Pepper, de Witt, and Colton, 1956; Colton and de Witt, 1958; de Witt and Colton, 1959; de Witt, 1960), we were more or less continuously entangled with the old "classic" nomenclature which was introduced in the fourth geological district by James Hall (1839, p. 322) and in the third geological district by Lardner Vanuxem (1842, p. 172). Not long after we began to study the rocks of the West Falls Formation—the rocks above the Sonvea Formation and below the Java Formation (Pepper, de Witt, and Colton, 1956)—we realized that the many usages of the name Portage as proposed by geologists during the past 125 years and the correlation of the units involved with the name were a source of confusion and difficulty. At different times, the name Portage had been applied to a single lithologic unit (Luther, 1897), to a group of several lithologic units (Hall, 1840, p. 391; 1843, p. 224), to a series of rocks in a time sense (Tilton and others, 1927, p. 88-90), and in parts of Kentucky, New York, Pennsylvania, and West Virginia to a magnafacies that included most of the marine Upper Devonian rocks. The name was given not only to a group of rocks exposed along the Genesee River in western New York but also to the upper massive sandstone unit within that group (Hall, 1843).

The relationship of the stratigraphic units that make up the Portage Group of Hall (1839) along the Genesee River in the fourth geological district and the units in the Portage or Nunda Group of Vanuxem (1842), about 75 miles to the east near Ithaca in the third geological district, illustrates some of the confusion surrounding the name Portage. Hall and Vanuxem believed the two groups were correlative.

In the type area along the Genesee River, Hall's Portage Group (1839) consisted of the strata above the Gardeau or Lower Fucoidal Group and below the Chemung Group. The Portage consisted of a thin basal sandstone, the Table Rock Sandstone of Chadwick (1933), a medial unit of intercalated shale and sandstone, and a thick upper massive sandstone. These units make up the upper part of our West Falls Formation at the Genesee. The upper massive sandstone is the Nunda Sandstone Member of our West Falls, and the two lower units make up the West Hill Member and the upper part of the Gardeau Shale Member of the West Falls.

In 1842 Vanuxem defined his Portage or Nunda Group near Ithaca as including the Sherburne Flagstone and Shale, the Cashaqua Shale, and the Gardeau and Portage Groups of the previous State reports. Vanuxem drew the top of his Portage Group at the base of the overlying Ithaca Group. Our study of these Upper Devonian rocks showed

that in the vicinity of Ithaca, the strata in the basal part of Vanuxem's Ithaca Group are the Renwick Shale Member, Ithaca Member, and West River Shale Member of our Genesee Formation. These strata are equivalent at the Genesee River to rocks which comprise the upper half of our Genesee Formation—the upper part of Hall's Genesee Slate. Thus, the strata near Ithaca that Vanuxem assigned to the Portage Group are not the lateral equivalents of Hall's Portage on the Genesee but are separated from the Portage beds by the lower part of the West Falls Formation and all the subjacent Sonyea Formation. The strata that compose the upper part of the West Falls Formation on the Genesee, Hall's Portage Group, change facies to the east; and if present in the Ithaca area, the equivalents of Hall's Portage Group occur in the upper part of the Chemung rocks well south and southwest of Ithaca. In their respective districts, both Hall and Vanuxem defined the top of their Portage Group largely by paleontologic criteria. They were not aware that because of changes in lithology the equivalents of Hall's original Portage Group lay within a dissimilar facies with a different fauna in the Ithaca area. If the original definition of the group (Hall, 1839) is strictly followed, none of Vanuxem's Portage at Ithaca is Portage because all the strata are older than the beds originally assigned to the group by Hall.

In 1843 Hall added to the problem by expanding his Portage Group on the Genesee to include the following units in descending order: Gardeau or Lower Fucoidal Group, Cashaqua Shale, and Sherburne Flagstones. The Sherburne Flagstones, the Sherburne Flagstone Member of our Genesee Formation, feathers out of the sequence to the east of the Genesee River in the area between Seneca and Cayuga Lakes and is unrepresented by coarse-grained rocks on the Genesee. As redefined by Hall, his expanded Portage Group includes the upper half of our Genesee Formation and all the younger Sonyea and West Falls Formations.

In the years between 1843 and 1943, the use of Portage Group, Portage Formation, or Portage Series, spread widely across the Appalachian basin as attempts were made by many geologists to synthesize the Upper Devonian stratigraphy of this area. Correlations were proposed partly on the basis of similar lithology and stratigraphic position but more commonly on the presence of similarity of faunas. Usage of the name Portage became so varied that each author was more or less obliged to redefine the unit in each paper. The wide latitude of usage of Portage increased the misunderstanding surrounding the unit. A glance through the papers of the New York Geological Survey and particularly the writing of G. H. Chadwick will illustrate many of the changes in the usage of Portage during the past 125 years.

From these and much other data, we concluded that the name Portage was largely invalid because of the multitudinous usages in the past. Furthermore, within the framework of stratigraphic nomenclature which we have proposed for the Upper Devonian rocks in the western half of New York, Portage need not be recognized in a formal sense. Consequently, we abandoned use of Portage rather than continue to compound the confusion surrounding the name by attempting to rigorously redefine the stratigraphic unit. The limitation on the length of text accompanying oil and gas charts (Pepper and others, 1956; Colton and de Witt, 1958) precluded a lengthy discussion of the stratigraphic nomenclature and possible revision; also, some of the most cogent reasons in favor of abandoning Portage as a group or formation were not clearly evident until we had completed work on the Sonyea Formation and the Genesee Formation.

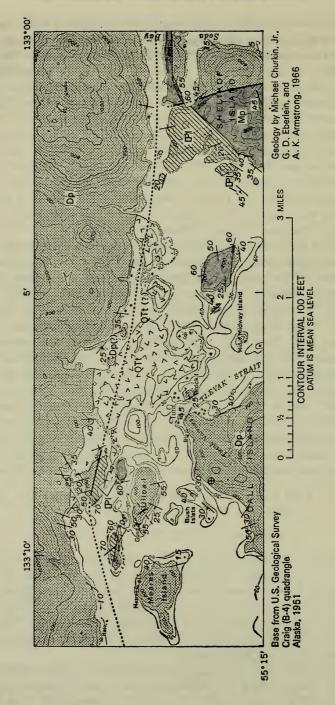
The New York Geological Survey did not include the Portage among the stratigraphic units recognized in the Upper Devonian sequence in the text of the most recent issue of the State geologic map (New York State Museum and Science Service Geological Survey, 1962). The New York Survey included the rocks that made up Hall's expanded Portage Group (Hall, 1843) as parts of our Genesee, Sonyea, and West Falls Formations. On the map, our formations were given group status. Clearly the action of the New York Geological Survey indicates the intent to abandon the Portage. Similarly Rickard's correlation chart of the Devonian rocks in New York (Rickard, 1964) does not contain the Portage as a recognized stratigraphic unit. Rickard shows in the chart that the "Chagrin" ("Portage") is a depositional phase—a gray shale, siltstone, limestone, and calcareous nodule facies—which is present at different places in several of the groups of strata in the Upper Devonian sequence in western New York.

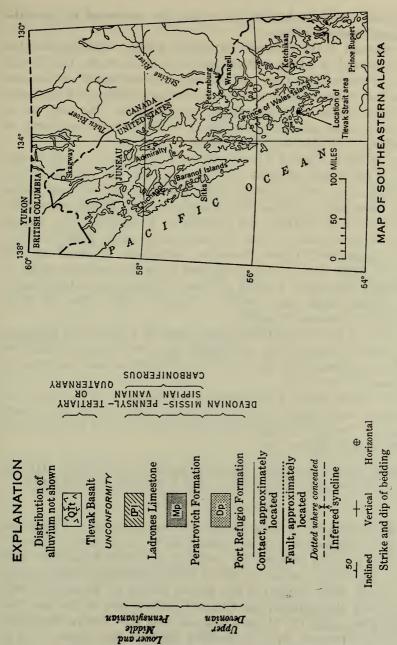
TLEVAK BASALT, WEST COAST OF PRINCE OF WALES ISLAND, SOUTHEASTERN ALASKA

By G. Donald Eberlein and Michael Churkin, Jr.

A number of small islands off the west coast of Prince of Wales Island at the latitude of Ketchikan (fig. 1) are underlain by one or more basalt flows. These flows are best exposed along the shoreline. The flows make up a new formation, here named the Tlevak Basalt for its occurrence in the vicinity of Tlevak Strait (fig. 1). The basalt is probably flat lying, and the low relief of the shoreline exposes only a few feet of lava at any locality.

Generally the basalt is strongly jointed and forms thin subhorizontal slabs. In places there are well-developed closely spaced concentric fractures along which the rock weathers spheroidally. Locally, the





Lower and Upper Mississiphan

FIGURE 1.—The Tlevak Strait area, southeastern Alaska.

basalt has vertical columnar joints that are transected by the sub-horizontal fracture system.

In contrast to the Paleozoic formations which predominate in this region, several of which contain basaltic volcanic rocks, the Tlevak Basalt is remarkably fresh and contains phenocrysts of olivine and plagioclase (labradorite) in an intergranular groundmass of labradorite microlites, olivine, and subordinate clinopyroxene. The rock has a normative mode and may be classified as olivine basalt according to the scheme of Yoder and Tilley (1962).

On the south shore of a small unnamed island about 1 mile north-west of the north entrance to Tlevak Strait, the Tlevak Basalt unconformably overlies steeply dipping beds of the Ladrones Limestone. Elsewhere, in discontinuous exposures, the Tlevak apparently overlies the Port Refugio and Peratrovich Formations. The Tlevak Basalt thus appears to rest unconformably upon Pennsylvanian, Mississippian, and Upper Devonian formations. A small area of similarly fresh olivine basalt is exposed along the northeast shore of Trocadero Bay about 9 miles northeast of the area of this report. This olivine basalt may overlie a formation older than the Port Refugio, the Descon Formation.

The top of the Tlevak Basalt is not exposed, and its total thickness is uncertain. Judging from its inferred gross subhorizontal structure and the maximum relief adjacent to shoreline exposures, it is probably less than 100 feet thick.

The Tlevak Basalt is demonstrably younger than the Ladrones Limestone and therefore is post-Pennsylvanian. It is tentatively assigned a Tertiary or Quaternary age because of its freshness and general lack of deformation.

FORMATIONS OF THE BISBEE GROUP, EMPIRE MOUNTAINS QUADRANGLE, PIMA COUNTY, ARIZONA

By Tommy L. Finnell

A thick sequence of shale, sandstone, conglomerate, and limestone of Early Cretaceous age is widespread in southeastern Arizona. Outcrops of these sedimentary rocks in the Mule Mountains near Bisbee (fig. 2) were called the Bisbee beds by Dumble (1902, p. 706). Later, they were designated the Bisbee Group by Ransome (1904, p. 56), who distinguished the Glance Conglomerate at the base, the Morita Formation, the Mural Limestone, and the Cintura Formation at the top. However, in areas north of the Mule Mountains, such as the Tombstone Hills and the Dragoon Mountains (Gilluly, 1956, p. 74), and to the northwest in the Mustang Mountains (Hayes and Raup, 1968), no lithologic equivalent of the Mural Limestone occurs, and the Bisbee was mapped as a formation with a basal conglomerate

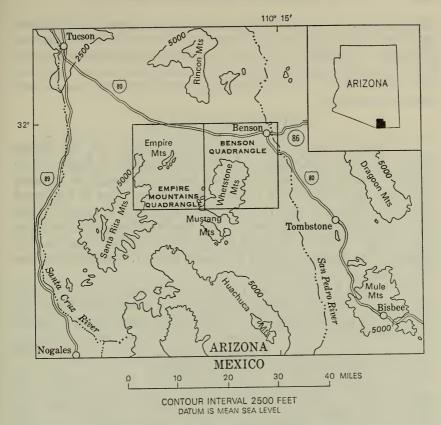


FIGURE 2.—Empire Mountains area, Arizona.

member. Cretaceous strata in the Whetstone Mountains, called Bisbee Group by Tyrrell (1957; 1965) and Bisbee(?) Formation by Creasey (1967), are partly correlative with the Bisbee Group in its type area (Haves and Drewes, 1968, p. 55); but, in the absence of an equivalent to the Mural Limestone, Tyrrell divided the group into the Willow Canyon Formation at the base, the Apache Canyon Formation, the Shelleburg Canyon Formation, and the Turney Ranch Formation at the top. Schafroth (1968, p. 60-64) recognized the same formations in the Empire and northern Santa Rita Mountains. Recent detailed studies confirm that the formations recognized by Tyrrell are also valid units in the Empire Mountains quadrangle, and the names are therefore adopted for use as described below. However, both earlier workers included the Glance Conglomerate as a member of the Willow Canyon Formation, whereas I have chosen to map the Glance as a separate formation. Hence, the Willow Canyon Formation in the Empire Mountains quadrangle is restricted to

those beds that occur stratigraphically above the Glance Conglomerate and below the Apache Canyon Formation. Also, the Shelleburg Canyon Formation of Tyrrell (1957) is here called Shellenberger Canyon Formation so that the spelling will be the same as that shown on the 1958 edition of the U.S. Geological Survey topographic map of the Benson quadrangle.

GLANCE CONGLOMERATE

The Glance Conglomerate, the basal formation of the Bisbee Group, consists of as much as 5,600 feet of pebble to boulder conglomerate and sedimentary breccia of variable lithology. In general, two types of conglomerate occur in the Glance. One is composed mainly of limestone and dolomite fragments and is present along the south, east, and north flanks of the Empire Mountains; the other is composed mainly of quartzite and granitoid fragments and is only present in the northern part of the range where it conformably overlies the limestone conglomerate.

The limestone conglomerate ranges in thickness from 1 to 400 feet and consists of closely packed subangular to subrounded fragments of limestone and dolomite as much as 3 feet across set in a matrix of reddish-brown to gray calcareous sandstone and sandy siltstone. In places, tabular blocks of Paleozoic formations as much as 1,000 feet across are embedded in the conglomerate.

The quartzite-granitoid conglomerate is as much as 5,200 feet thick, and it consists of closely packed angular to subangular fragments of quartzite and gneissic quartz diorite as much as 3 feet across set in a matrix of reddish-brown to greenish-gray sandstone and sandy siltstone. The quartzite fragments are most numerous at the base, giving way gradually upward to the quartz diorite fragments.

The Glance rests on an erosional unconformity of widely variable relief. Along the south and east flanks of the Empire Mountains, the surface of unconformity has irregularities measured in tens of feet that are reflected in the variations of thickness of the conglomerate. In the northern part of the range, however, the unconformity was cut over a mass of Precambrian granitoid and metamorphic rocks that stood several thousand feet higher than the floor of the basin of Bisbee deposition. There, the Glance consists of debris eroded from the highland and deposited along the basin margin. This Glance grades southward by intertonguing into the Willow Canyon, Apache Canyon, and the lower 1,000 feet of the Shellenberger Canyon Formations; all of them are part of a conformable sequence above the limestone conglomerate to the south. Therefore, the upper contact of the Glance in the southern part of the range is placed at the top

of the highest conglomerate bed that is overlain by siltstone or sandstone that contains only a few thin beds of conglomerate, but in the area of intertonguing to the north, tongues of conglomerate are shown as Glance where they exceed 20 feet in thickness; otherwise, such tongues are arbitrarily not mapped separately.

WILLOW CANYON FORMATION

The Willow Canyon Formation was named by Tyrrell (1957) for exposures along Willow Canyon on the west flank of the Whetstone Mountains, and he described the type section in the SE¼ sec. 25, T. 18 S., R. 18 E. The upper member of Tyrrell's Willow Canyon Formation in the type section is similar to the upper 1,000 feet of the Willow Canyon Formation (restricted) in the Empire Mountains, where the formation is at least 3,000 feet thick. Because the lithology of the Willow Canyon Formation is more varied in the Empire Mountains, reference section there is designated that extends from the S½ sec. 18 to the NW¼ sec. 29, T. 18 S. R., 17 E. The reference section contains the lower two-thirds of the formation on the northwest side of a zone of normal strike faults and the upper three-fourths on the southeast side, and it shows the typical lithologic types of the formation.

The Willow Canyon Formation in the reference section consists predominantly of an alternating sequence of sandstone and siltstone that erodes to valleys and ridges of moderate relief. The sandstone is light yellowish gray, light pinkish gray, and light yellowish brown, arkosic, crossbedded, and locally conglomeratic. The pebbles and cobbles in the conglomeratic sandstone are well rounded and composed mainly of quartzite and some granodiorite and quartz diorite. The siltstone is commonly dark reddish brown in the lower two-thirds of the formation and olive gray to greenish gray in the upper one-third. A few thin beds of silty limestone and calcareous sandstone occur in the upper 300 feet of the formation, and a shaly siltstone about 250 feet below the top contains fossil gastropods of the genus *Viviparus* (W. A. Cobban, written commun., 1966), a long-ranging fresh-water form.

The contact with the overlying Apache Canyon Formation is placed arbitrarily at the horizon above which limestone is a dominant part of the lithology. Thus, some thin limestone beds are included in the Willow Canyon Formation. The Willow Canyon is at least 3,000 feet thick in the southern part of the Empire Mountains, but northward it grades laterally into the Glance Conglomerate by intertonguing and is absent along the north side of the range.

APACHE CANYON FORMATION

The Apache Canyon Formation was named by Tyrrell (1957) for exposures in Apache Canyon on the west flank of the Whetstone Mountains. He described the type section at that locality extending from the base in the SW¼NW¼ sec. 31, T. 18 S., R. 19 E., to the top in the NW¼NW¼ sec. 1, T. 19 S., R. 18 E.

The Apache Canvon Formation in the Empire Mountains and the western part of the Whetstone Mountains consists of thinly laminated to thick-bedded silty limestone, black shale, siltstone, and arkosic sandstone in an alternating sequence that erodes to low-rounded ridges and valleys. The limestone is generally dark gray to black and exhibits laminae as little as a millimeter thick of black limestone alternating with gray silty limestone that forms a distinctive lithologic type unlike any Paleozoic limestone in the region. The shales are dark-gray to black fissile shales intercalated with the limestone beds. Siltstone is dark gray to orange and red, calcareous, and thin bedded to massive. Massive red siltstone about 200 feet above the base contains a bed of gypsum about 2 feet thick, and some limestone beds above the gypsum contain numerous siltstone casts of what may have been gypsum crystals. The upper half of the Apache Canyon contains several beds of sandstone as much as 15 feet thick. The sandstone is yellowish gray, yellowish brown, and olive gray, fine to very coarse grained, partly crossbedded and partly massive, and arkosic. Biscuit-shaped bodies of silty limestone as much as 3 feet across and half a foot thick occur here and there in the limestone and shale beds.

The Apache Canyon Formation grades upward through increasing amounts of sandstone and siltstone into the Shellenberger Canyon Formation, the upper contact in the Empire Mountains being placed at the top of the uppermost limestone that is more than 3 feet thick and that is overlain by siltstone and sandstone containing only a few beds of limestone. The Apache Canyon Formation is at least 1,600 feet thick in the Empire Mountains but only 550 to 740 feet thick in the type area (Tyrrell, 1957, p. 101). Like the Willow Canyon Formation it intertongues northward with the Glance.

SHELLENBERGER CANYON FORMATION

The Shellenberger Canyon Formation was named by Tyrrell (1957) for exposures along the northeast-trending part of Shellenberger Canyon on the southwest flank of the Whetstone Mountains. Tyrrell (1957) described the type section of the Shellenberger Canyon Formation as extending from the base in the NE½NE½ sec. 34, T. 18 S., R. 18 E., to the top in the SW½NE½ sec. 4, T. 19 S., R. 18 E.

The Shellenberger Canyon Formation in the Empire and Whetstone Mountains consists of an alternating sequence of shale, siltstone. sandstone, and a little limestone that erodes to valleys and ridges of moderate relief. The shale and siltstone are commonly shades of olive brown, olive gray, and greenish gray although in a few places they are reddish brown. Small bits of carbonized plant fragments are present in many of the somber-hued shales and siltstones, and silicified logs as much as 5 feet in diameter abound in some beds. The sandstone is olive brown, olive green, olive gray, and pinkish gray, fine to very coarse grained, arkosic, massive to crossbedded and lenticular. Some of the sandstone beds are conglomeratic and contain pebbles of quartzite and chert as much as an inch in diameter. One sandstone bed about 500 to 800 feet below the top of the formation has a distinctive basal conglomerate as much as a foot thick that contains well-rounded pebbles of black and red chert. The siltstone above this sandstone contains a calcareous zone about 2 to 10 feet thick that contains scattered well-rounded pebbles and cobbles of limestone and quartzite.

Limestone is mainly restricted to the lower 1,300 feet of the formation; some of it is laminated like that in the Apache Canyon Formation and some is thin bedded. Two beds about 1,000 and 1,300 feet above the base are composed of pelecypod and gastropod shells in a matrix of calcite, sand, and silt. Fossils of an Early Cretaceous dinosaur (Miller, 1964, p. 378) were obtained by Moore and Miller (1960, p. 59-60) from a shaly bed between the shell beds in the Empire Mountains, and a fossil clam *Trigonia* n. sp., was found in calcareous siltstone that overlies a similar shell bed in the Whetstone Mountains (Tyrrell, 1957).

The contact of the Shellenberger Canyon Formation with the overlying Turney Ranch Formation in both ranges is placed at the top of a light-brown sandstone that is overlain by red shale and siltstone that contains nodules of gray limestone and that is overlain in turn by light yellowish-gray to light pinkish-gray sandstone. The Shellenberger Canyon generally ranges in thickness from 3,950 to 4,330 feet, but to the north, the lower 1,000 feet of the formation grades into the Glance Conglomerate.

TURNEY RANCH FORMATION

The Turney Ranch Formation was named by Tyrrell (1957) for its typical exposures immediately north and west of Turney Ranch in the Whetstone Mountains. This ranch is now known as the Clyne Ranch and is located in secs. 21, 22, and 27, T. 19 S., R. 18 E., just inside the southeast corner of the Empire Mountains quadrangle. The

described type section extends from the base in the SE¼NW¼ sec. 11, T. 19 S., R. 18 E., to the top in the SW¼ sec. 15, T. 19 S., R. 18 E.

The Turney Ranch Formation in the Empire and Whetstone Mountains consists of red siltstone and shale alternating with light-pinkish-gray sandstone. It erodes to valleys and ridges of moderate relief. The siltstone and shale are various shades of red and purplish red though locally gray near nodules of limestone. Some beds are calcareous. In places, gradations from red fine-grained sandstone into siltstone and finally into shale occur both vertically and laterally. The sandstone is light pinkish gray to pale orange, medium to coarse grained, arkosic, and crossbedded. Lenses of pebble conglomerate containing chert and quartzite pebbles set in an arkosic matrix occur here and there in the sandstone beds. Scour-and-fill structures a few inches deep occur at the base of some of the sandstone beds.

In the type area, the Turney Ranch Formation is overlain with angular unconformity by upper Cenozoic gravel deposits. In the Empire Mountains, tilted and faulted Turney Ranch beds are locally overlain with angular discordance by conglomerate and andesitic breccia that are probably of Late Cretaceous age, for they are very similar to the lower part of the Salero Formation which has been assigned a Late Cretaceous age by Drewes (1968) on the basis of potassium-argon age determinations of the upper part of the formation. The Turney Ranch Formation is at least 3,200 feet thick in the type area. Its original thickness cannot be determined because its top is eroded.

AGE

The formations of the Bisbee Group in the Empire and Whetstone Mountains are all part of a conformable sequence and are considered to be of the same general age. The only fossils that are diagnostic as to age are of an Early Cretaceous dinosaur and of the mollusk Trigonia, all from the Shellenberger Canyon Formation near the middle of the sequence. Tyrrell (1957) submitted specimens of Trigonia to Prof. A. A. Stovanow, who determined that they are a new species that is not represented in the fauna from the Bisbee Group in the type area. However, Stoyanow considered that the beds containing Trigonia in the Whetstone Mountains are probably no older than the upper Trinity Group (Upper Cretaceous) and that they were deposited in the same sea as the Bisbee Group in the Bisbee area. The Bisbee Group in the Empire Mountains quadrangle rests unconformably on a variety of older rocks, the youngest of them being sedimentary and volcanic rocks that are probably equivalent to the Gardner Canvon Formation of Triassic age (Drewes, 1968). The Bisbee is overlain with a strong angular unconformity by sedimentary and volcanic rocks that are similar to the Salero Formation of Late Cretaceous

age. These geologic relations, together with the meager fossil evidence, indicate that the Bisbee Group in this area was probably deposited during Early Cretaceous time.

PANTANO FORMATION

BY TOMMY L. FINNELL

The Pantano Formation, herein adopted, was introduced by C. F. Tolman in unpublished data on the geology of the Tucson 30-minute quadrangle prepared for the U.S. Geological Survey about 1912. The first published detailed description of the Pantano, other than a brief reference by King (1939, p. 1692), was by Brennan (1962, p. 46–53). In this report Brennan selected and described the type section along the south side of U.S. Highway 80 (now Interstate 10) extending from about the east line of sec. 2, T. 17 S., R. 17 E. (top of section), westward along the highway to the center of the W½ sec. 31, T. 16 S., R. 17 E., in the Empire Mountains quadrangle, southeastern Arizona. Detailed mapping has revealed that several normal faults repeat parts of the Pantano Formation in Brennan's type section. Also at least three angular unconformities within the formation omit certain stratigraphic units from the type section.

The Pantano Formation is readily divisible into five units that

have a combined thickness of at least 6,400 feet:

Unit 1, the lowest unit, is faulted against vertical beds of the Bisbee Formation of Early Cretaceous age along Interstate 10, but it lies on eroded edges of Tertiary or Cretaceous volcanic and sedimentary rocks in Barrel Canyon near State Highway 83 about 11 miles south-southwest of the town of Pantano. It consists of reddish-brown bouldery mudstone and siltstone, conglomeratic sandstone, conglomerate, and at least one rhyolite tuff of probable ash-flow origin. The sandstone and conglomerate are mostly in poorly defined layers, but in places they form alternating beds as much as 2 feet thick. This unit is at least 1,600 feet thick.

Unit 2 is red and yellowish-gray mudstone alternating with fine conglomerate and conglomeratic sandstone of the same colors; in the upper 100 feet it contains at least four beds of olive-gray partly oolitic sandy argillaceous limestone. The whole unit is about 500 feet thick.

Unit 3 is massive to poorly bedded reddish-brown conglomerate with a few mudstone and coarse-grained sandstone partings. It pinches out against the lower units a few hundred feet south of the type section but extends northward for at least 2 miles. It is at least 2,500 feet thick.

Unit 4 is a flow of dark-purplish-gray andesite called Turkey-Track Porphyry by Cooper (1961) because of the large plagioclase phenocrysts. The upper part of the flow is eroded, and it thins to a featheredge about 300 feet south of the highway.

Unit 5 locally contains a dark-purplish-gray volcanic conglomerate at the base. Where present, this grades up into purplish-gray and yellowish-brown conglomerate and coarse conglomeratic sandstone that grade upward and northward into fine-grained tuffaceous sandstone, siltstone and gypsiferous claystone containing a few interbeds of sandstone, and angular conglomerate or monolithologic breccia.

At least three tuff beds occur near the middle. The unit is at least 1,800 feet thick.

The subangular to well-rounded pebbles, cobbles, and boulders in the Pantano Formation were mainly derived from the Bisbee Formation, but partly from the Precambrian, Paleozoic, lower Tertiary(?), and Upper Cretaceous rocks of the region. In unit 3, boulders and cobbles of rhyodacite welded tuff identical in appearance to an Upper Cretaceous welded tuff are a conspicuous but minor constituent.

The surface upon which the Pantano Formation was deposited had considerable relief, as indicated by apparent depositional thinning and pinch-out of the lower units. Local deformation during deposition is suggested by angular unconformities between some of the units. For example, unit 5 rests on the andesite flow (unit 4) along the highway, but about half a mile to the south, it rests unconformably on limestone of unit 2.

In places, faulted and tilted beds of the Pantano are overlain by generally flat-lying gravel deposits of Pliocene(?) and Pleistocene age. The Pliocene(?) gravels are light pinkish gray to light yellowish gray, and they may have covered the entire type area of the Pantano Formation at some earlier time for remnants of these gravels are known to overlie the upper four units.

No indigenous fossils have been found in the Pantano Formation, although its Paleozoic and Lower Cretaceous clasts are fossiliferous. However, its age can be deduced from radiometric dating. Dates on minerals from the rhyolite tuff of unit 1 have been determined by Damon and Bikerman (1964, p. 69) as follows:

Sanidine (K-Ar) 36.7 ± 1.1 m.y.

Biotite (K-Ar) 32.8 ± 2.7 m.y.

P. E. Damon (written commun. 1966) determined the radiometric age of plagioclase from the andesite flow to be 24.4±2.6 m.y. (K-Ar). On this basis, the Pantano Formation seems to range in age from early Oligocene to early Miocene.

BARDSTOWN MEMBER OF THE DRAKES FORMATION IN CENTRAL KENTUCKY

By Warren L. Peterson

Prepared in cooperation with the Kentucky Geological Survey

The Drakes Formation of Late Ordovician age consists of two members at its type locality in south-central Kentucky: the Rowland Member and the overlying Preachersville Member (Weir and others, 1965). In the vicinity of Bardstown, Ky., the Drakes Formation is

divided into three members: the lowest, the Rowland Member, is closely similar to the Rowland in south-central Kentucky; the highest is the Saluda Dolomite Member; the middle member is herein named the Bardstown Member for exposures in the vicinity of Bardstown, Nelson County, Ky. The type section (fig. 3) is in Nelson County, Ky., in the southwest quarter of the Maud quadrangle in a roadcut along U.S. Highway 150, 1.35 miles northwest of Fredericktown, Washington County, Ky., and 0.95 mile from the west border of the quadrangle (lat 37°46′20″ N.; long 85°21′30″ W.).

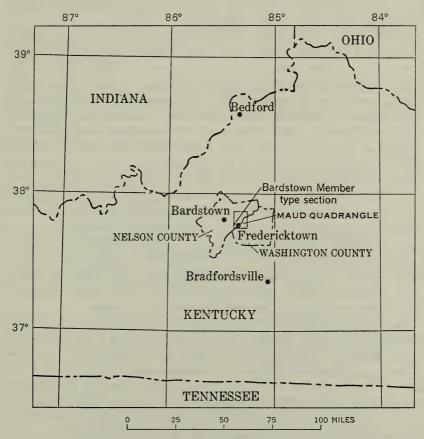


FIGURE 3.—Location of the type section of the Bardstown Member of the Drakes Formation.

About 90 percent of the Bardstown Member consists of fine- to medium-grained, gray to greenish-gray muddy limestone that contains scattered to abundant fossils and fossil fragments. Weathered rock is yellowish gray with a shaly aspect. Samples of the matrix yielded 27 to 50 percent insoluble residue. Beds are discontinuous and lentic-

ular and commonly 1 to 8 inches thick. About 10 percent of the member is gray bioclastic and coquinoidal limestone composed of whole fossils and fossil fragments in a fine-grained to very coarse grained matrix cemented in part by sparry calcite. The matrix is predominantly calcitic fossil debris, with variable amounts of non-carbonate silt and clay. It occurs commonly in ½- to 3-inch-thick discontinuous beds, lenses, and knots with rough to smooth surfaces that form thin ledges in gully exposures or slump as resistant blocks. A small percentage of the member is calcareous shale (containing more than 50 percent insoluble residue) that is physically indistinguishable from the muddy limestone.

Fossils in the Bardstown Member are principally brachiopods, bryozoans, pelecypods, gastropods, horn corals, and colonial corals. Horn corals ("Streptelasma") are sparse to abundant and, in the vicinity of Bardstown, are probably restricted to this member. The colonial corals are Tetradium (with tiny rectangular corallites) and Foerstephyllum, Favistella, and Caleopecia (with larger hexagonal corallites) (Browne, 1964). The corals with hexagonal corallites occur in heads as much as 4 feet across. The colonial coral heads are concentrated in two to four layers in the middle two-thirds of the unit. Locally, layers as much as 4 feet thick are 25 to 50 percent coral heads. The coral layers in the Bardstown Member are probably correlatives of the Otter Creek coral bed, which occurs at the base of the Preachersville Member of the Drakes Formation in eastcentral Kentucky (Simmons and Oliver, 1967). Massive coral layers in some exposures were called the "Bardstown coral reef" by Foerste (1903). Fossils in the Bardstown Member have been listed by Butts (1915), Browne (1964, 1965), and Hatfield (1968) under different formation names discussed in another part of this paper.

The top of the Bardstown Member is placed where the muddy fossiliferous limestone grades abruptly into dolomite of the overlying Saluda Dolomite Member. The contact is not obvious but can be located by application of dilute hydrochloric acid. The contact at the base of the Bardstown Member is commonly sharp or gradational through a few inches. In some places, however, the basal contact is gradational by interlayering through an interval up to 4 feet thick. The contact is placed at the base of rock with abundant fossils.

The Bardstown Member is commonly 25 to 35 feet thick but ranges in thickness from 12 to 40 feet within a 15-mile radius of the type section. It is recognizable along the outcrop northward to near Bedford, Trimble County, Ky., and southward and southeastward at least as far as Bradfordsville, Marion County, Ky.

The Bardstown Member has previously been called the unnamed member of the Drakes Formation by Peterson (1966, 1967, 1968).

7.3

20.6

Approximately the same unit was designated the Liberty Limestone by Nosow (1959, fig. 10), the Liberty Formation by Butts (1915), the Liberty Formation and lower part of the Whitewater Formation by Browne (1964), and part of the Tanners Creek Formation by Hatfield (1968). The Liberty Limestone (or Formation) which was named in Indiana (Nickels, 1903) has been recognized in Kentucky only on a faunal basis, and the name is, therefore, considered inappropriate for a rock-stratigraphic unit in Kentucky. The name Tanners Creek Formation defined by Fox (1962) at Tanners Creek, Dearborn County, Ind., was rejected for use in Indiana by Brown and Lineback (1966) because of the lack of a regionally traceable base. The name is not considered usable in Kentucky for the same reason. The rocks of the Bardstown Member have been considered to be of Richmond (Late Ordovician) age by all writers.

Drakes Formation at type section of the Bardstown Member of the Drakes Formation

[Section measured in the southwest quarter of the Maud quadrangle, Nelson County, in roadcut along U.S. Highway 150, 1.35 miles northeast of Fredericktown, Washington County, Ky., lat 37°46′20″ N.; long 85°21′30″ W.]

Thickness (feet) 22. Dolomite, light-gray to light-brown, fine- to medium-grained__ 4+ Drakes Formation: Saluda Dolomite Member: 21. Shale, grayish-green, silty and clayey; contains no megafossils; poorly exposed______ 2, 3-2. 8 20. Dolomite similar to unit 19 with abundant Tetradium heads; in part recrystallized, flattened parallel to 19. Dolomite, greenish-gray; color banded parallel to bedding on surface into 1 to 2 inch bands of yellowish brown, reddish brown, dark gray, and grayish green; banding probably reflects faint bedding; very fine grained; slightly silty and clayey; unfossiliferous except for poorly preserved dolomitized cylindrical bryozoans; some recrystallized calcitic bryozoans in upper 2 feet; dolomite pebbles in upper 2 feet; worm bored throughout; forms massive, jutting rounded ledge_____ 6.3 18. Dolomite, greenish-gray, very fine grained, massive; less resistant than overlying dolomite, silty and clayey;

Bardstown Member:

Brassfield Dolomite (basal part only):

17. Limestone, greenish-gray, muddy, fine-grained; contains scattered to abundant fossils and fossil fragments; in faint beds about 2 inches thick; worm bored; contains sparse thin beds, lenses, and knots of purer, more re-

Total Saluda Dolomite Member____

contains no megafossils; worm bored in lower 3 feet; upper 3 to 6 inches slightly indented and less resistant__

Drakes Formation—Continued	Thickness
Bardstown Member—Continued sistant gray fossiliferous limestone; fossils include brachiopods, cylindrical bryozoans, gastropods, pelecy-	(feet)
pods, and horn coral	7. 6
occur closely packed in parts of unit and sparsely in other parts	4. 3
15. Limestone, greenish-gray, muddy, fine-grained; contains sparse to abundant fossils and fossil fragments; in beds 1 to 5 inches thick; less abundant but more resistant limestone makes up about 20 percent of rock, is medium gray, has medium- to coarse-grained matrix and abundant fossils and fossil fragments in beds, lenses, and knots as much as 2 inches thick; minor olive-gray shale occurs interbedded with limestone; all gradations occur from nearly pure limestone to calcareous shale. Fossils include brachiopods, cylindrical bryozoans, horn corals,	1.0
gastropods, pelecypods, and scattered colonial coral	7. 2
heads 14. Limestone, greenish-gray, muddy; contains scattered	1. 2
fossils; colonial coral heads as much as 2 feet in diameter	1005
sparse to abundant 13. Limestone and minor shale; similar to unit 15	
12. Limestone and shale; limestone is greenish gray, contains patches of unidentified dark-green mineral, is muddy, fine to medium grained, in beds 1 to 6 inches thick with undulating surfaces; fossils generally sparse; lowest bed contains brachiopods and bryozoans; small gastropods found near top. Shale is olive gray, calcareous, weakly fissile; contains no megafossils, in partings and beds as much as 1½ inches thick. Rock in unit has similarities	
to both Bardstown and Rowland lithologies	3. 0
Total Bardstown Member	29. 3
Rowland Member: 11. Limestone, greenish-gray, medium grained; abundant small fossil fragments; single bed faintly parted into 1-inch layers	1. 5
10. Limestone, greenish-gray to grayish-green, very fine grained, silty and clayey, dolomitic, thinly laminated; parted into 1- to 2-inch layers; contains no megafossils; weakly resistant; weathers with shaly aspect; some mud-	
9. Limestone, greenish-gray; similar to unit 10 but contains scattered to abundant fossil fragments and ostracode shells; in beds 1 to 11 inches thick; no conspicuous lami-	16. 8
nations; forms resistant outcropping ledge	5. 9
8. Limestone; similar to unit 10	12. 2
contains no megafossils	. 3

Drakes Formation—Continued	TI to to
Rowland Member—Continued	Thickness: (feet)
6. Limestone, greenish-gray, impure, very fine grained, laminated; parts into ½- to 1½-inch layers; contains some fossil fragments in upper part; more resistant to weath-	V ,
ering than units such as 10	4. 9
5. Limestone similar to unit 4 in beds as much as 1 inch thick	
and limestone similar to unit 3 mostly as thin partings	1. 0
4. Limestone, light-olive-gray, fine- to medium-grained; contains pebbles of greenish-gray limestone; single bed	
with smooth top and bottom; contains no megafossils	. 3
3. Limestone, grayish-green; conspicuously green on outcrop, silty and clayey, very fine grained; contains sparse small brachiopods, bryozoans, and fossil fragments; weathers	
shaly	1. 8
2. Shale, dark-olive-gray, calcareous, clayey, silty, weakly	
fissile; contains no megafossils	. 3
Total Rowland Member	45. 0
Grant Lake Limestone:	
1. Limestone and shale (not measured).	

THE CANTWELL FORMATION OF THE CENTRAL ALASKA RANGE

By Jack A. Wolfe and Clyde Wahrhaftig

NAME AND TYPE AREA

The Cantwell Formation was first described by Eldridge (1900, p. 16) as the Cantwell Conglomerate from exposures in the canyon of the Nenana River, then called the Cantwell River. Its type locality is considered to be the east wall of the canyon of the Nenana River from the mouth of Slime Creek northward for about 7 or 8 miles in the Healy C-4 quadrangle (fig. 4, loc. 1). This type locality agrees with Eldridge's original text (1900, p. 1) but not with his map (map 3) nor with the statement in Wahrhaftig (1958, p. 8), both of which are in error.

DISTRIBUTION

The Cantwell Formation occupies a large synclinorium extending along the center of the Alaska Range from the headwaters of the Wood River westward to the Muldrow Glacier (fig. 4). Brooks and Prindle (in Brooks, 1911) and Reed (1961) have mapped bodies of the Cantwell farther southwest along the north base of the range. Capps (1933) mapped a large area of the Cantwell Formation on the south side of the Denali fault between Foggy Pass and Anderson Pass, an area separated from the main body of the Cantwell by a belt of Paleozoic and Triassic rocks along the north side of the fault.

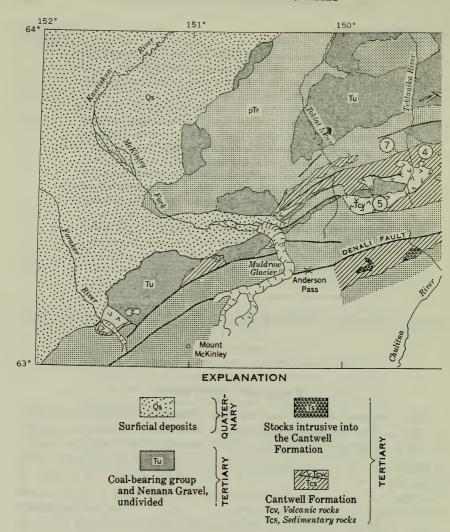
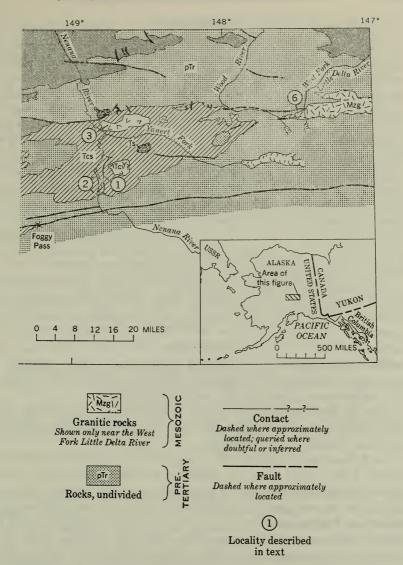


FIGURE 4.—Geologic sketch map of the central Alaska Range showing distributext. Geology modified from Reed (1961, pl. 1),



tion of the Cantwell Formation, related formations, and localities referred to in Wahrhaftig (1958, pl. 1), and Capps (1912, pl. 2).

LITHOLOGY

The Cantwell Formation consists predominantly of interbedded conglomerate, sandstone, argillite, shale, and coal but also contains volcanic rocks, especially near its top. The Cantwell is intruded by an abundance of sills and dikes ranging in composition from diabase to rhyolite and by monzonite stocks as much as 3 to 4 miles across.

At two localities the Cantwell Formation is clearly younger than large intrusive bodies.(1) At the pass between the headwaters of the Wood River and the West Fork Little Delta River (fig. 4, loc. 6), it rests unconformably on granodiorite and quartz monzonite at the west end of a large batholith which was traced eastward by Capps (1912, pl. 2) as far as Delta Creek, a distance of 30 miles; the granodiorite beneath the contact has been weathered to grus for a thickness of 20 feet.(2) On the east side of the Muldrow Glacier, dikes believed to be feeders to the volcanic rocks of the Cantwell Formation cut the intrusive body underlying Mount Eielson (Reed, 1933).

The Cantwell Formation is generally moderately well consolidated and is locally very well indurated. Some beds in the formation are extremely well sorted pebble conglomerate consisting of quartz, chert, quartzite, and argillite pebbles with little or no matrix. Few of the pebbles are less than half, or more than twice, the median size. The pebbles are indented and moulded against each other, possibly as a result of solution and redeposition of silica under tectonic pressure; the conglomerate therefore has negligible porosity.

Coal in the Cantwell Formation is generally of bituminous rank, even where beds are tightly folded and the conglomerate compressed, as in the vicinity of mafic dikes.

Dark flows and rhyolite tuffs occupy an open syncline centered on Mount Fellows (fig. 4, loc. 3). A belt of predominantly silicic volcanic rocks is exposed at the top of the formation from Double Mountain (fig. 4, loc. 4) westward to the Muldrow Glacier. These silicic volcanic rocks give Polychrome Pass (fig. 4, loc. 5) its color variety. In both the Mount Fellows and Polychrome Pass areas, the volcanic rocks are at the top of the section, yet only 1,000 to 3,000 feet of sedimentary rocks lies between the volcanic rocks and the base of the formation. The volcanic rocks at Mount Fellows and Polychrome Pass may be equivalent to lower parts of the Cantwell Formation at other places, or they may have erupted in areas where highlands in the pre-Cantwell topography rose a few thousand feet above the base of the formation elsewhere.

The lithology of the clasts in the Cantwell Formation varies from place to place, suggesting that the formation was derived from at least three different sources. In the type area along the Nenana River, the sandstone and conglomerate are predominantly dark gray and con-

sist largely of argillite, chert, quartzite, and quartz grains and pebbles. The source of this dark-gray facies was probably south of the Alaska Range, possibly in the Mesozoic rocks in the Talkeetna Mountains or the southern Alaska Range. Along the north edge of the area of outcrop of the Cantwell east of the Nenana River, this facies interfingers with light-brown to white sandstone and conglomerate consisting largely of schist fragments derived from the Birch Creek and Totatlanika Schists immediately to the north. From the vicinity of Polychrome Pass (fig. 4, loc. 5) westward, the dark-colored Cantwell is replaced by light-brown conglomerate and sandstone that consist largely of clasts of gray limestone in a light-brown matrix of unknown composition and origin. The westward limit of this light-colored limestone-bearing facies is unknown. Its source could have been the area of Devonian limestone near the crest of the range (and on the north side of the Denali fault) immediately south of its area of outcrop.

The limestone-bearing facies has not been mapped in the Cantwell Formation south of the Denali fault, as this body has not been visited since Capps' reconnaissance in 1930. A study of this facies and its correlation with facies north of the fault might give information on the lateral displacement of the fault since Paleocene time.

THICKNESS

The total thickness of the Cantwell Formation, whose upper surface has been eroded, is unknown. Generally, the thickness ranges from 2,000 to 5,000 feet. The maximum thickness is about 10,000 feet in a reference section (fig. 4, loc. 2) in the walls of the Nenana Canyon between Clear Creek and Carlo (Wahrhaftig, 1958, pl. 3) but may include an unknown thickness of sills.

AGE

The Cantwell Formation has had a varied history of age assignments. Eldridge (1900, p. 16), when first describing the Cantwell, was unable to assign an age to the formation. Brooks (1911, p. 78–83), who, with Prindle, mapped the Cantwell Formation from Mount McKinley to the Nenana River (fig. 4), correlated it on the basis of lithologic similarity with the Nation River Formation, then considered to be of Carboniferous age; but Brooks, in making this correlation had to explain away as faulted inliers beds containing plant remains which he had collected and which F. H. Knowlton reported to be Tertiary.

Moffit (1915) found plant fossils in rocks about 15 miles east of the Nenana that are similar to those of the type area. Knowlton and Hollick (in Moffit, 1915, p. 48) assigned these fossils to the Eocene but noted that Brooks' collections were similar. For 20 years there-

after, geologists working in Alaska were troubled by the fact that both the Cantwell Formation and an apparently much younger coal-bearing formation of the Nenana coal field to the north contained plants of

supposed Eocene age.

In 1936, Ralph Chaney made new collections in the Cantwell at another highly fossiliferous locality and determined the fossils to be Cretaceous in age (Chaney, 1937), the age for the Cantwell quoted by Capps (1940, p. 118). In a further refinement of its age, Imlay and Reeside (1954, p. 235), noting the similarity of the plant species identified from the Cantwell by Chaney to those in the Upper Cretaceous Chignik Formation of the Alaska Peninsula and the Lower Cretaceous Melozi and Kaltag Formations of western Alaska, placed the Cantwell in the Lower Cretaceous Albian Stage.

Fossil plants have been collected from several localities in the Cantwell Formation, but only one locality, that of Chaney's collections, has furnished well-preserved and abundant material. Reexamination of Chaney's collections (Univ. of California Museum Paleont. loc. P3654—loc. 7 of fig. 4) indicates that the following taxa are present:

Glyptostrobus? sp.

Metasequoia occidentalis (Newb.) Chan.

Sparganium antiquum (Newb.) Berry

Planera microphylla Newb.

Cocculus flabella (Newb.) Wolfe

Cissus sp. aff. C. marginata (Lesq.) R. W. Br.

Grewiopsis auriculaecordatus (Holl.) Wolfe

Dicotylophyllum flexuosa (Newb.) Wolfe

All these named species are indicators of a Tertiary age; Cissus marginata is known from both Upper Cretaceous and Paleocene rocks (Brown, 1962). The species of Planera, Cocculus, Grewiopsis, and Dicotylophyllum were recorded from the Chickaloon Formation of the Cook Inlet region (Wolfe and others, 1966) and from other rocks thought or known to be of Paleocene age. Inasmuch as the particular species of these four genera have not been found in rocks younger than Paleocene, we therefore consider the known fossil plants to be indicative of a Paleocene age for the Cantwell Formation.

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